

Dynamic and thermodynamic controls on midlatitude clouds: What can the observed satellite record teach us?

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National Center for Atmospheric Research

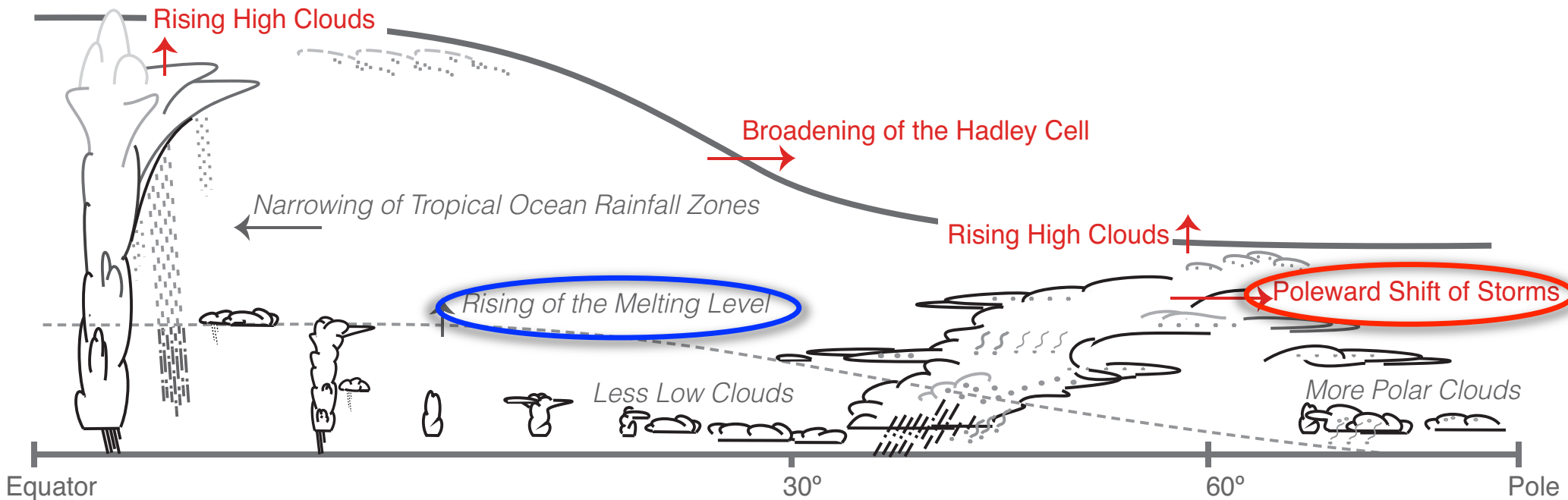
Rossow Symposium on Clouds, Their Properties, and Their Climate Feedbacks

New York, NY

June 6, 2017



Introduction



IPCC AR5 (Chapter 7)

Can we put observational constraints on the dynamic and thermodynamic processes underlying midlatitude cloud feedbacks?

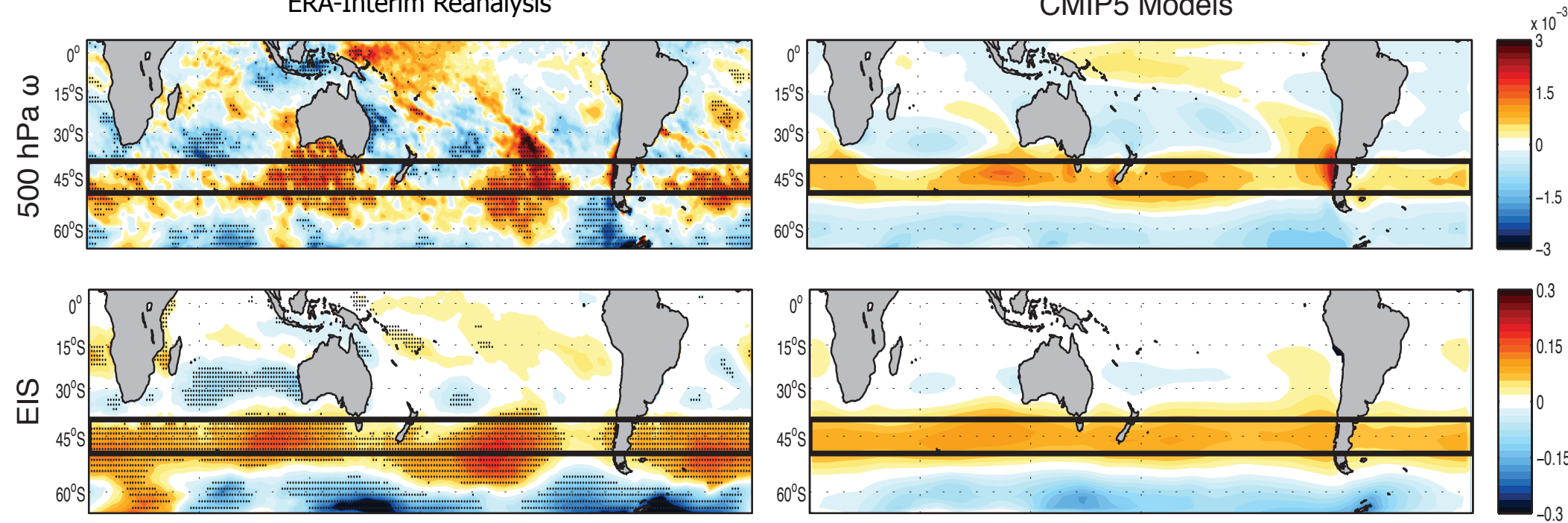
- 1) Positive cloud feedback with poleward shift of midlatitude jet stream
- 2) Negative cloud phase feedback at midlatitudes

Poleward Jet Shift: Dynamical Anomalies

1° Poleward Jet Shift (Interannual Variability)

ERA-Interim Reanalysis

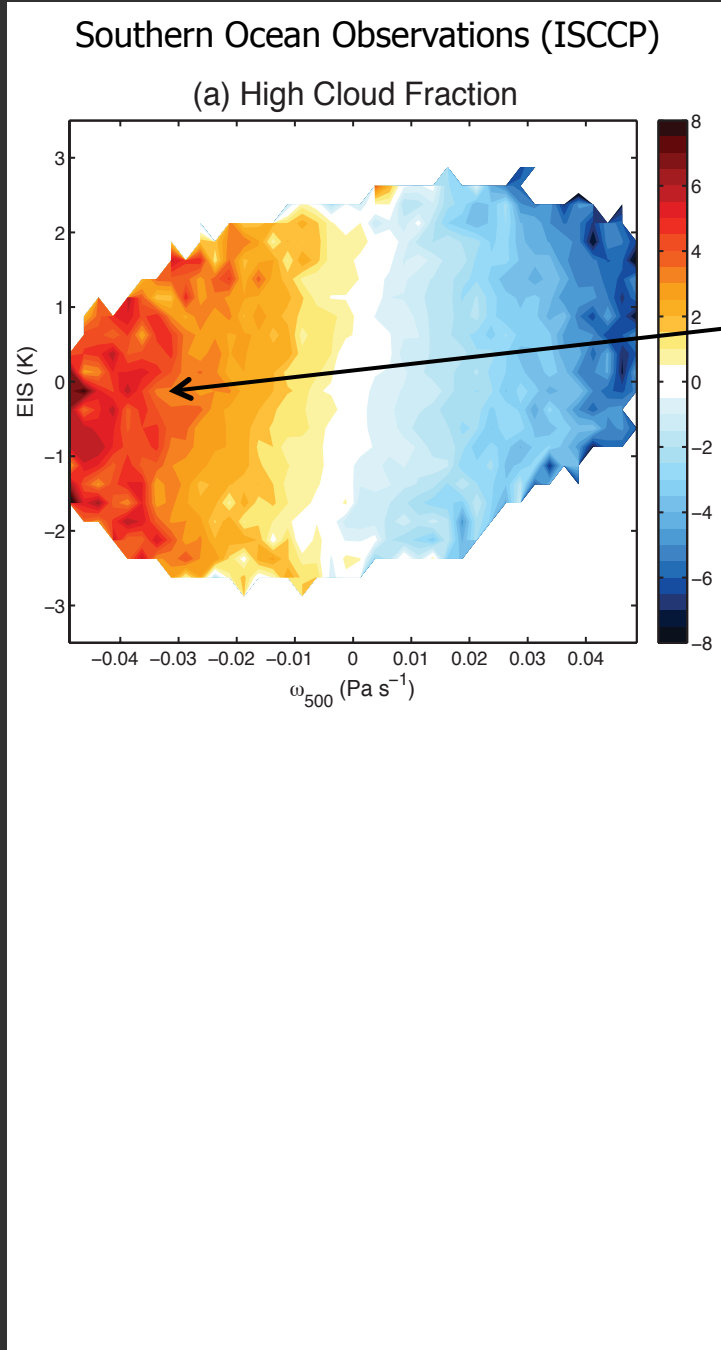
CMIP5 Models



Grise and Medeiros (2016)

When the jet shifts poleward, downward vertical velocity and increased lower tropospheric stability (EIS) anomalies occur equatorward of the jet.

Poleward Jet Shift: Cloud Controlling Factors

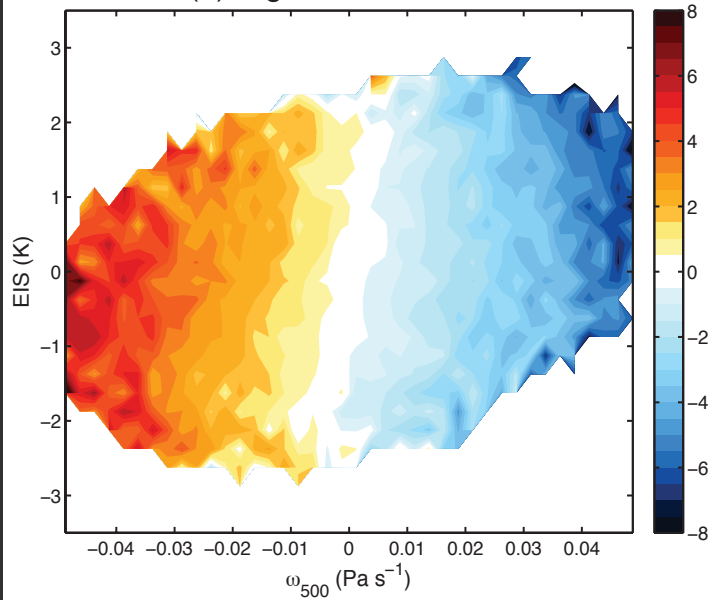


High clouds increase with upward vertical velocity anomalies.

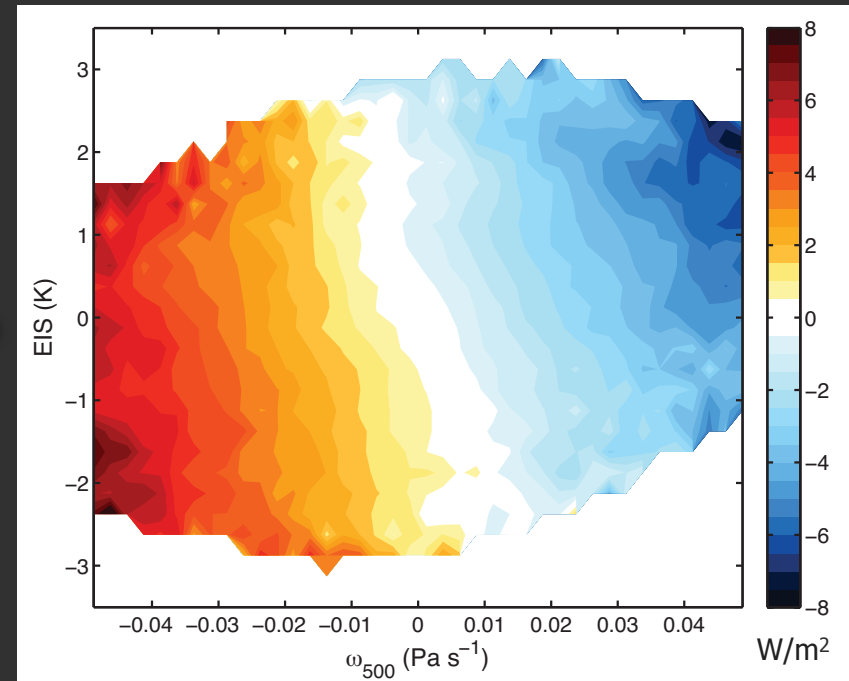
Poleward Jet Shift: Cloud Controlling Factors

Southern Ocean Observations (ISCCP)

(a) High Cloud Fraction



Longwave Cloud-Radiative Effect (CERES)

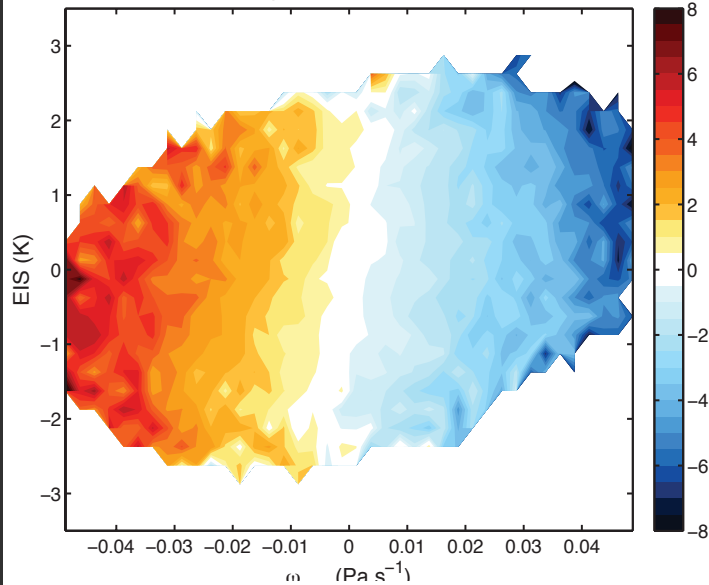


Longwave cloud-radiative effects also closely follow vertical motion.

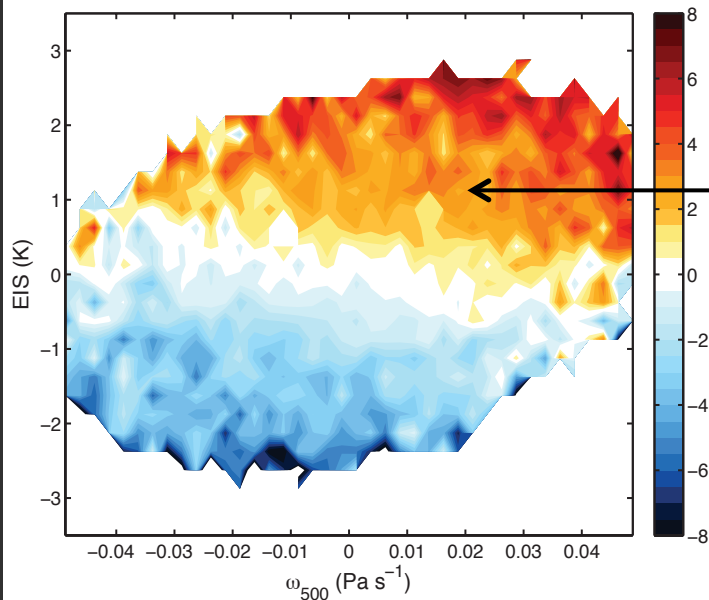
Poleward Jet Shift: Cloud Controlling Factors

Southern Ocean Observations (ISCCP)

(a) High Cloud Fraction



Low Cloud Fraction (Random Overlap)



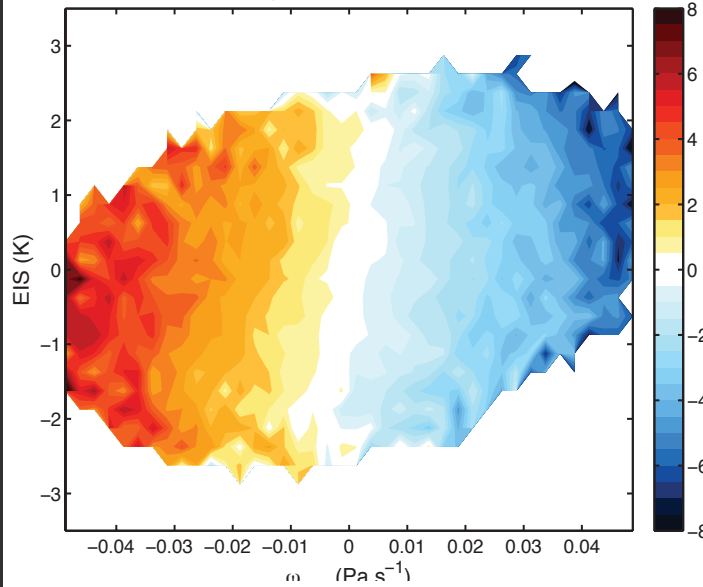
Low clouds increase with increasing strength of boundary layer inversion. (Wood and Bretherton 2006)

Poleward Jet Shift: Cloud Controlling Factors

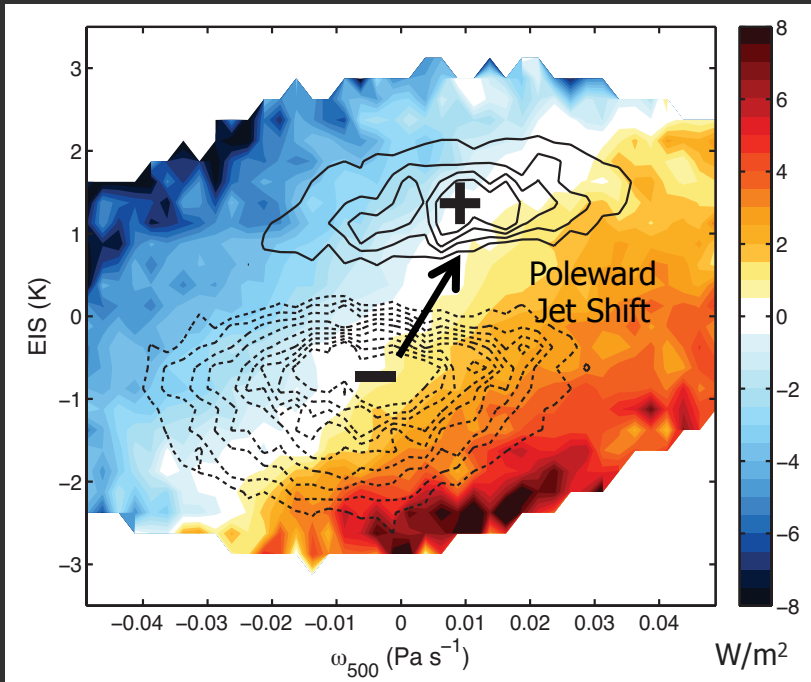
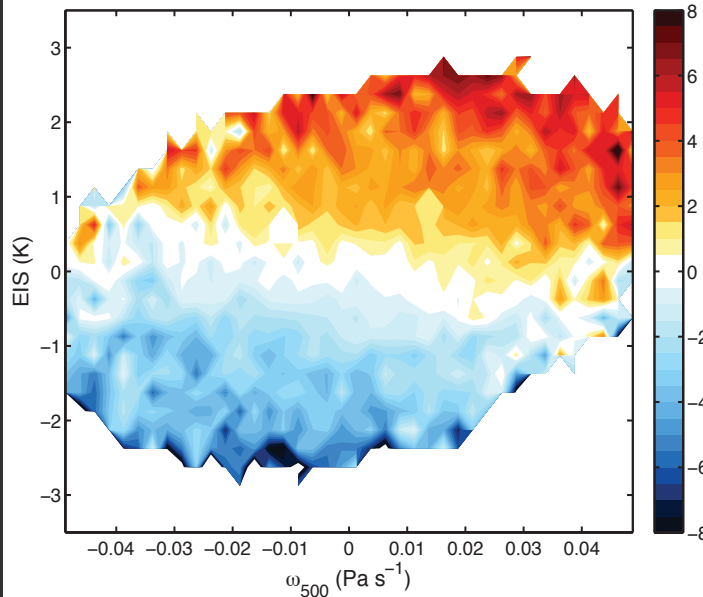
Shortwave Cloud-Radiative Effect (CERES)

Southern Ocean Observations (ISCCP)

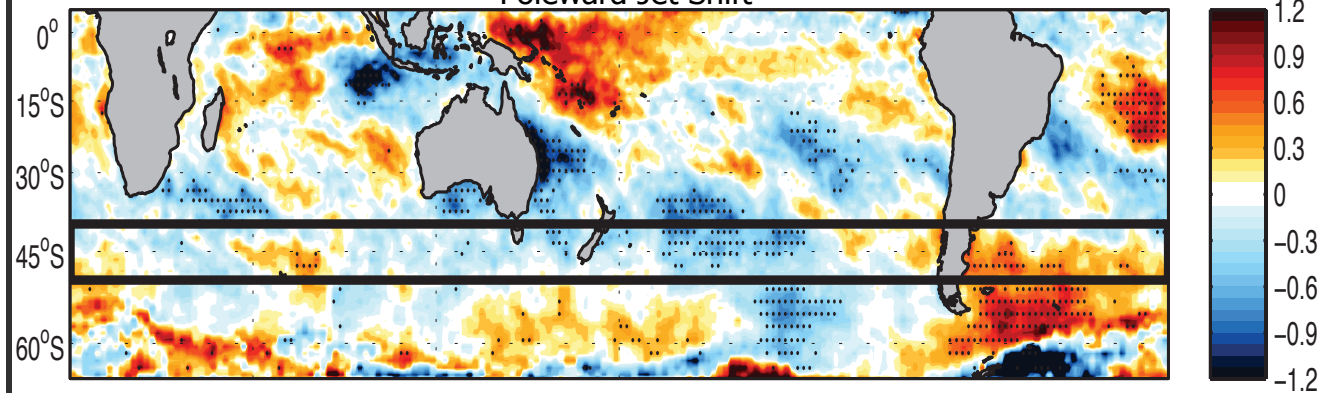
(a) High Cloud Fraction



Low Cloud Fraction (Random Overlap)



Poleward Jet Shift



Shortwave cloud-radiative effects depend on both vertical velocity and EIS anomalies.
(see also Myers and Norris 2013)

Poleward Jet Shift: Model Biases

Shortwave Cloud-Radiative Effect Anomalies
1° Poleward Jet Shift (Interannual Variability)

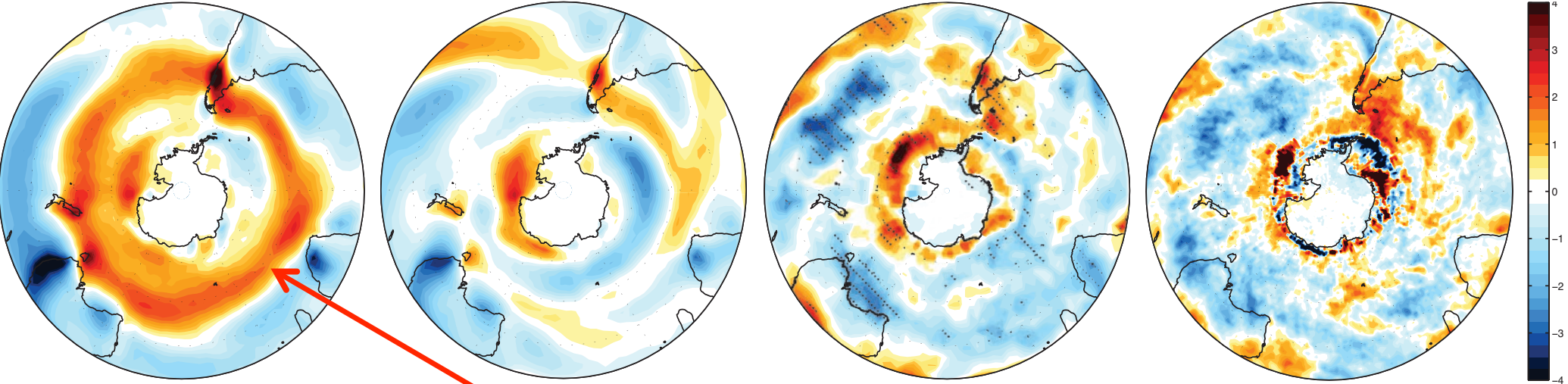
CMIP5 Models: "Type I"

CMIP5 Models: "Type II"

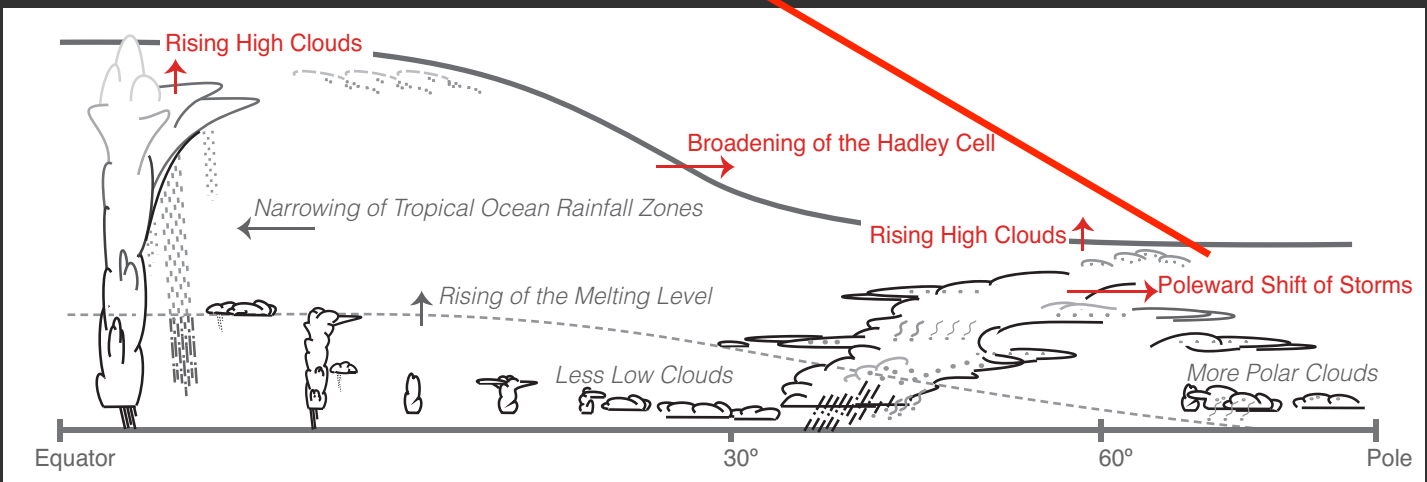
ISCCP-FD

CERES

W/m²



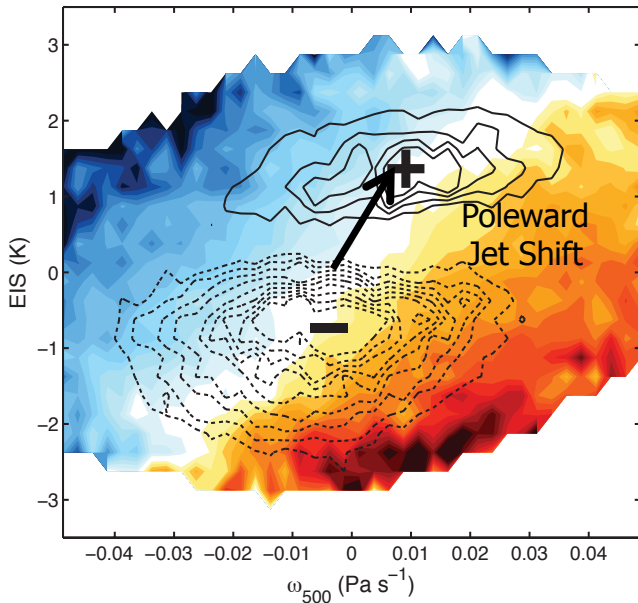
Grise and Polvani (2014)



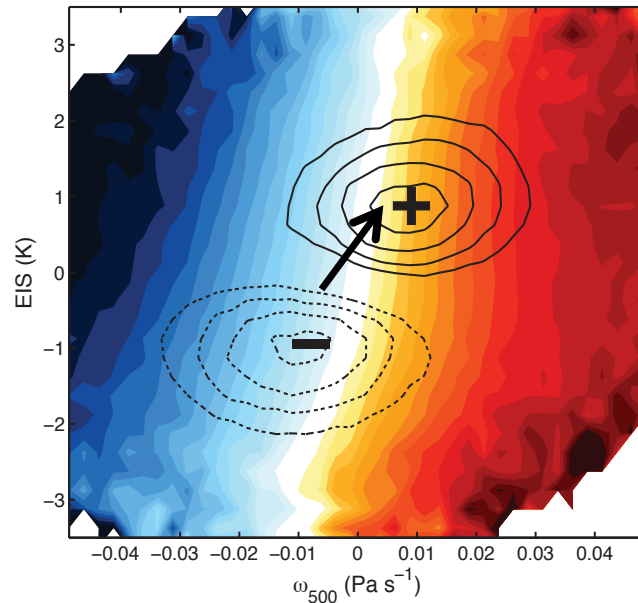
Poleward Jet Shift: Model Biases

Southern Ocean: Shortwave Cloud-Radiative Effect

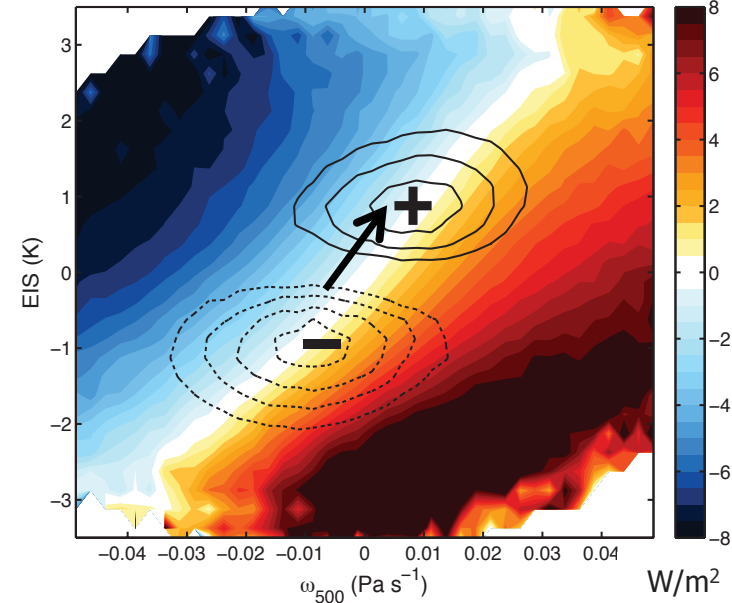
Observations (CERES)



CMIP5 Models: Type I



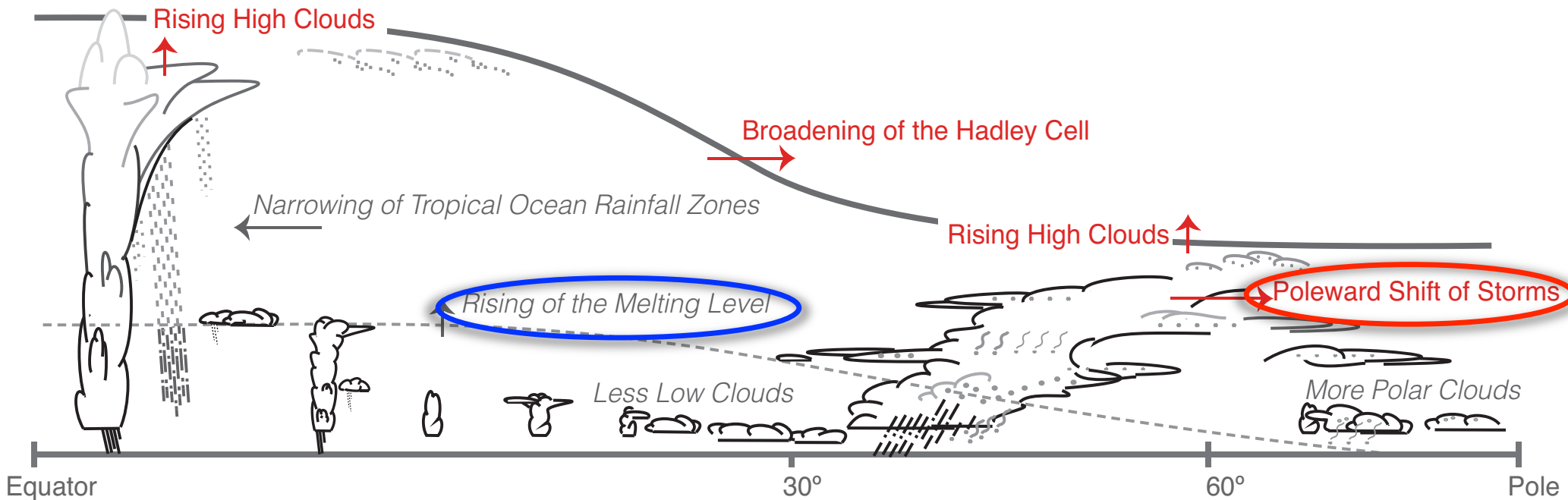
CMIP5 Models: Type II



Grise and Medeiros (2016)

Biases in type I models arise from an underestimation of the sensitivity of model cloud-radiative effects to the strength of the boundary layer temperature inversion (EIS).

Next: Cloud Phase Feedback



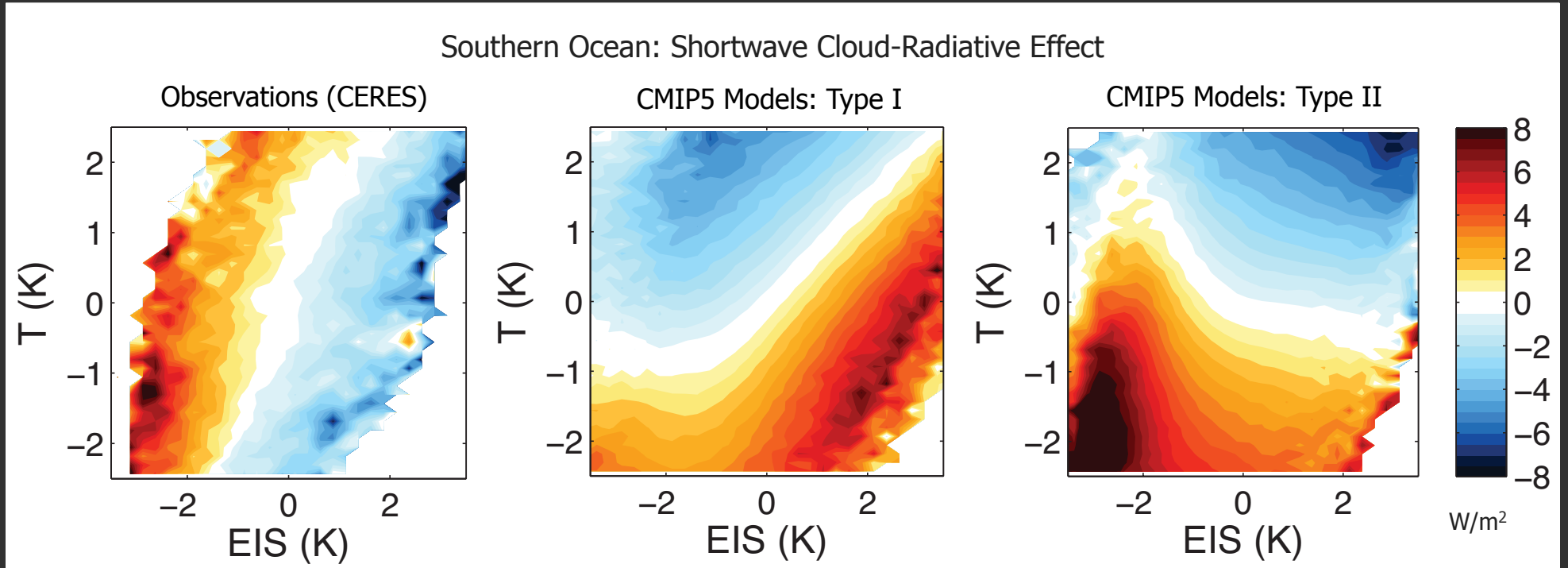
IPCC AR5 (Chapter 7)

Can we put observational constraints on the dynamic and thermodynamic processes underlying midlatitude cloud feedbacks?

- 1) Positive cloud feedback with poleward shift of midlatitude jet stream
- 2) Negative cloud phase feedback at midlatitudes

Cloud Phase Feedback: Controlling Factors

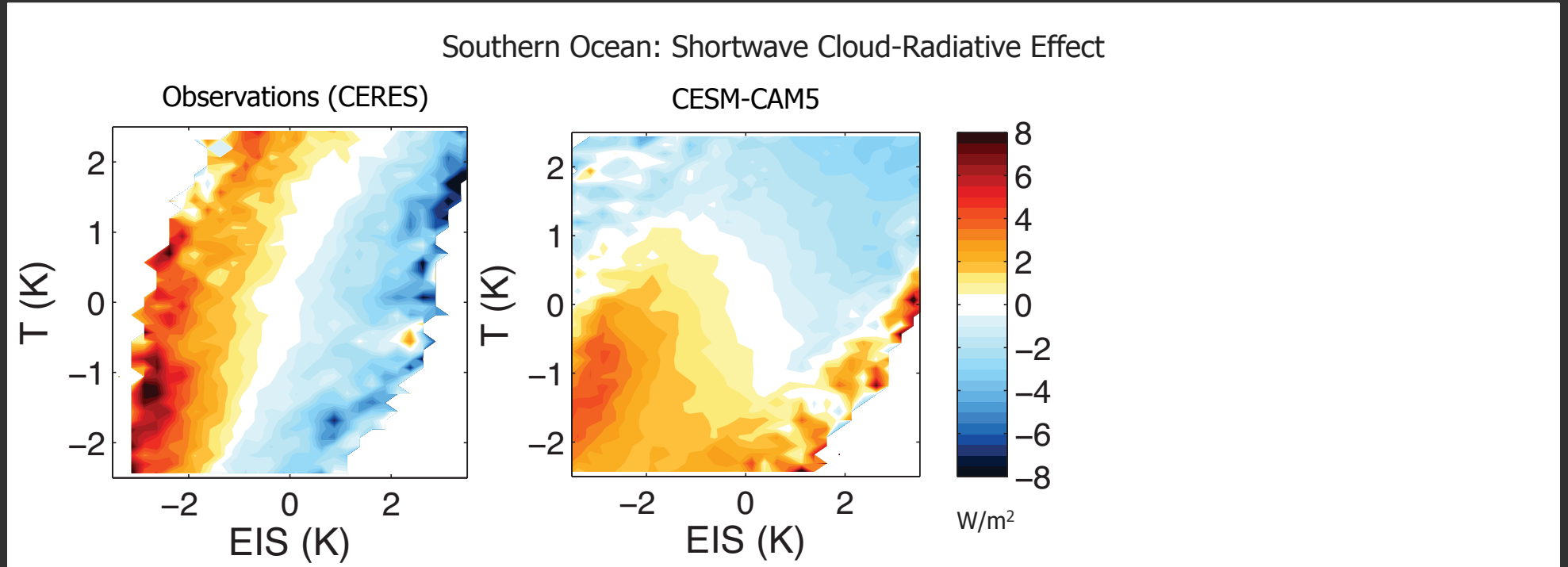
Consider two controlling factors on Southern Ocean shortwave cloud-radiative effects: EIS and temperature (500–850 hPa)



Shortwave cloud-radiative effect variability with temperature is overestimated by CMIP5 models (see also Terai et al. 2016).

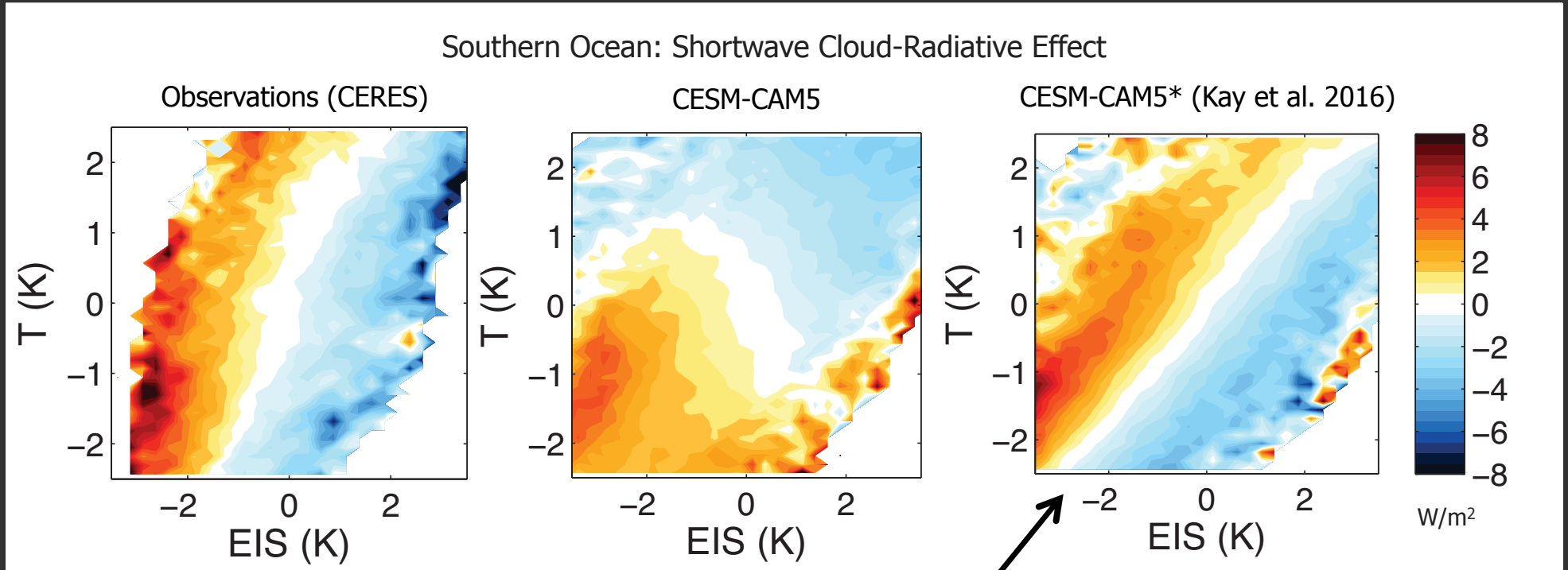
Cloud Phase Feedback: Controlling Factors

Consider two controlling factors on Southern Ocean shortwave cloud-radiative effects: EIS and temperature (500–850 hPa)



Cloud Phase Feedback: Controlling Factors

Consider two controlling factors on Southern Ocean shortwave cloud-radiative effects: EIS and temperature (500–850 hPa)



Kay et al. (2016) modified CESM-CAM5, increasing supercooled water at the expense of ice in mixed-phase clouds.

Summary

Can we put observational constraints on the dynamic and thermodynamic processes underlying midlatitude cloud feedbacks?

1) Positive cloud feedback with poleward midlatitude jet shift

Likely overestimated as many CMIP5 models underestimate the observed dependence of midlatitude clouds on the strength of the boundary layer temperature inversion (EIS)
(Grise and Medeiros 2016)

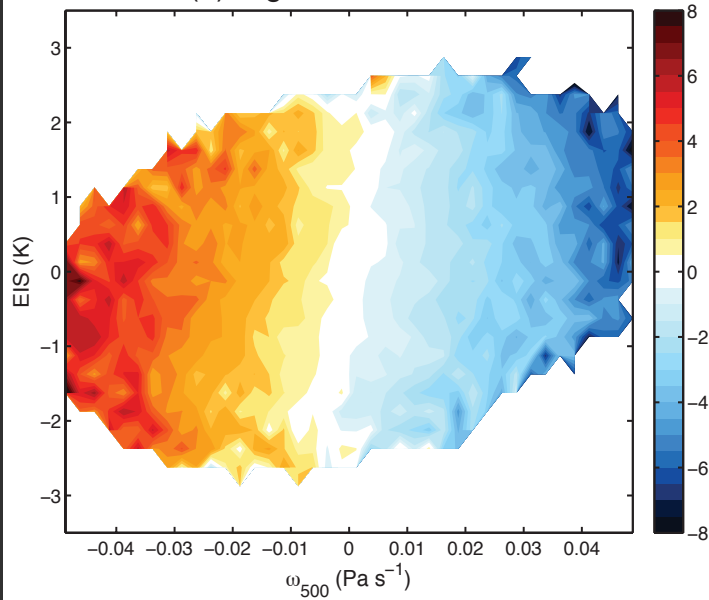
2) Negative cloud phase feedback at midlatitudes

Likely overestimated as CMIP5 models overestimate the observed dependence of midlatitude clouds on temperature variability (see also Terai et al. 2016, Tan et al. 2016, etc.)

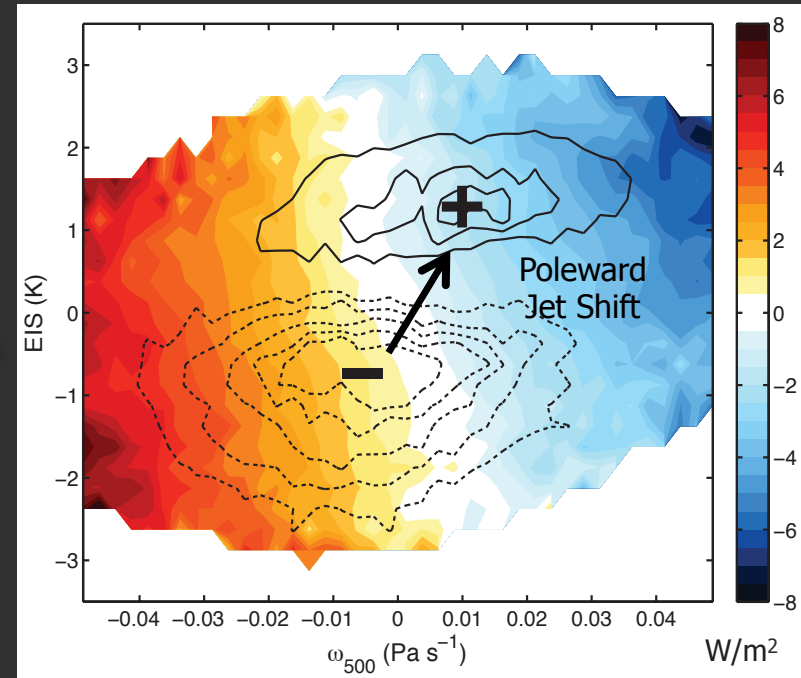
Poleward Jet Shift: Cloud Controlling Factors

Southern Ocean Observations (ISCCP)

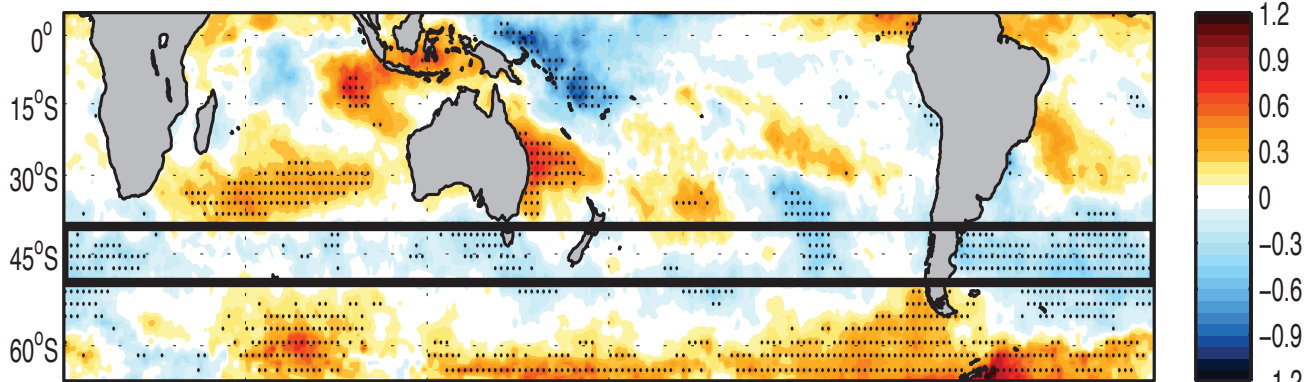
(a) High Cloud Fraction



Longwave Cloud-Radiative Effect (CERES)



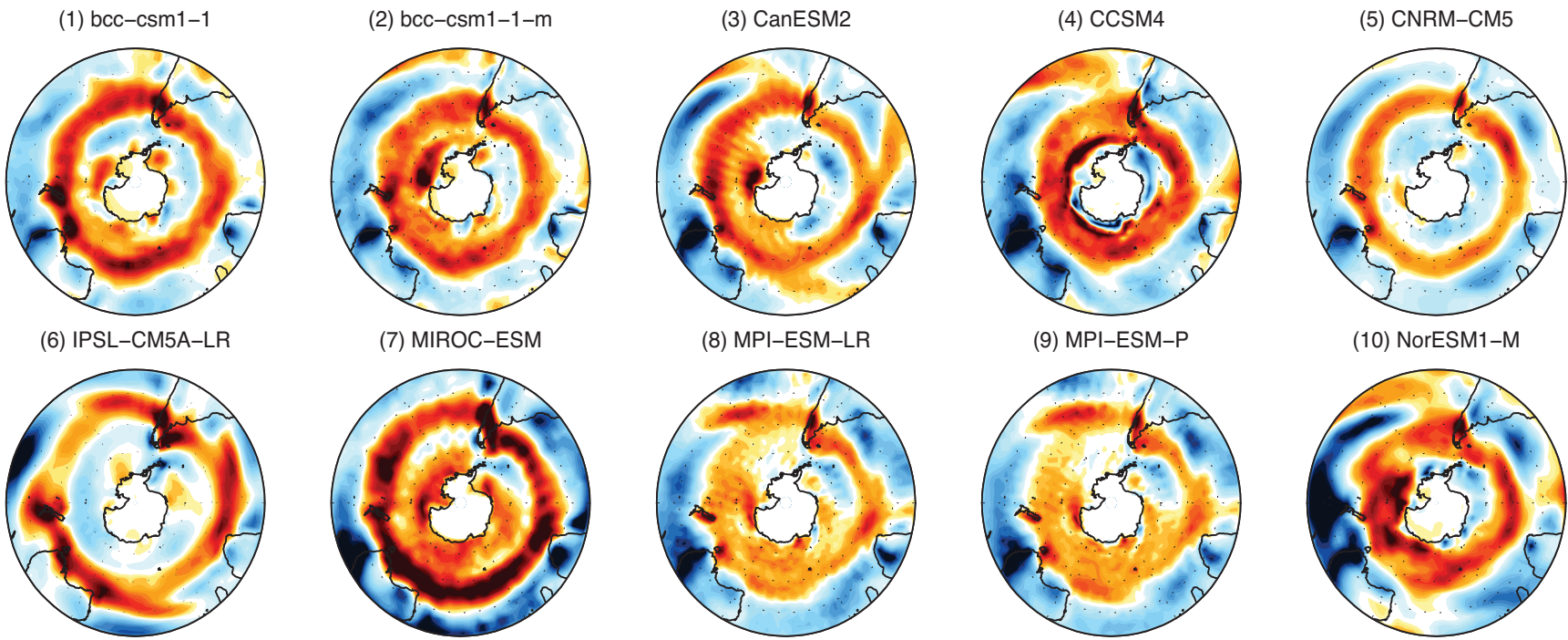
Poleward Jet Shift



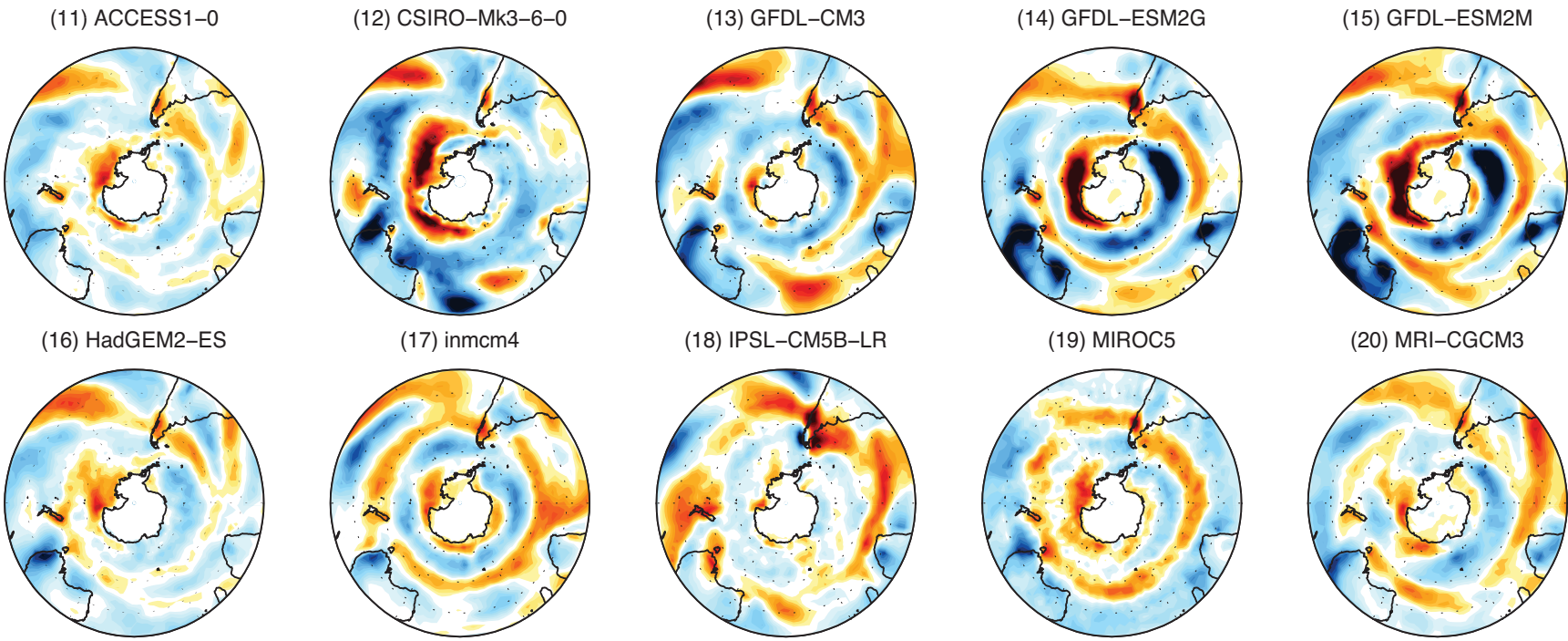
Longwave cloud-radiative effects closely follow vertical motion.

CMIP5 Models

Type I models



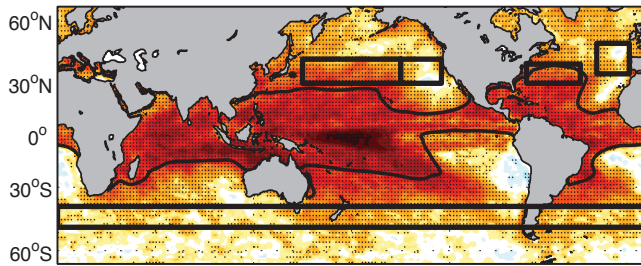
Type II models



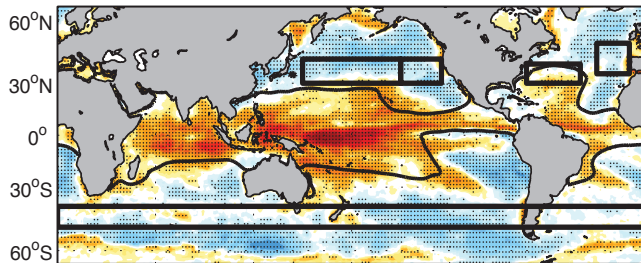
Understanding the Model Biases

Observations (CERES)

Correlations of SW CRE with ω_{500}

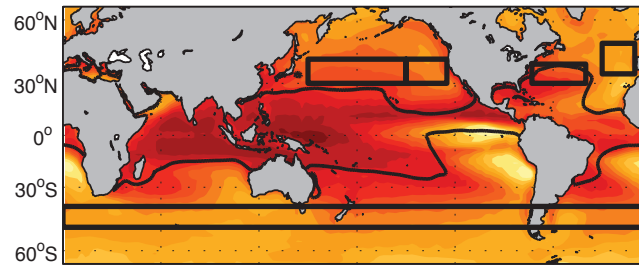


Correlations of SW CRE with EIS

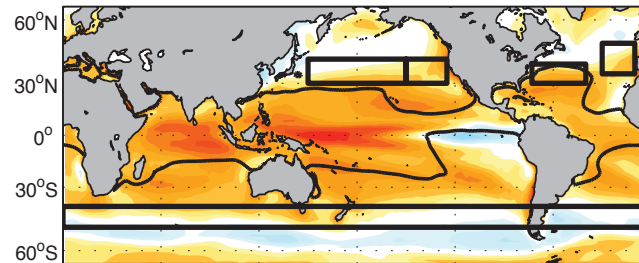


Type I Models

Correlations of SW CRE with ω_{500}

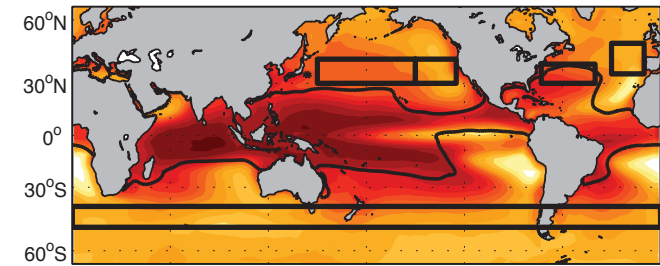


Correlations of SW CRE with EIS

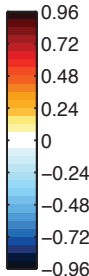
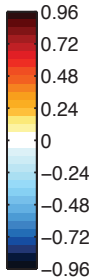
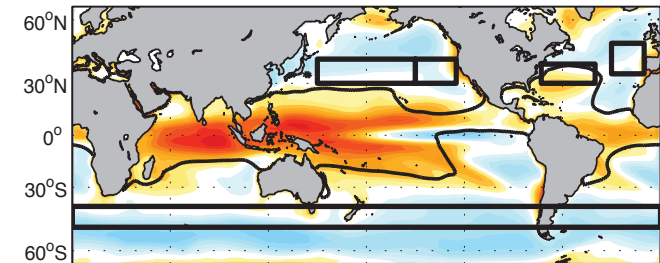


Type II Models

Correlations of SW CRE with ω_{500}



Correlations of SW CRE with EIS



Grise and Medeiros (2016)

Biases in type I models appear related to overdependence of model cloud-radiative effects on vertical motion and underdependence of model cloud-radiative effects on lower tropospheric stability.

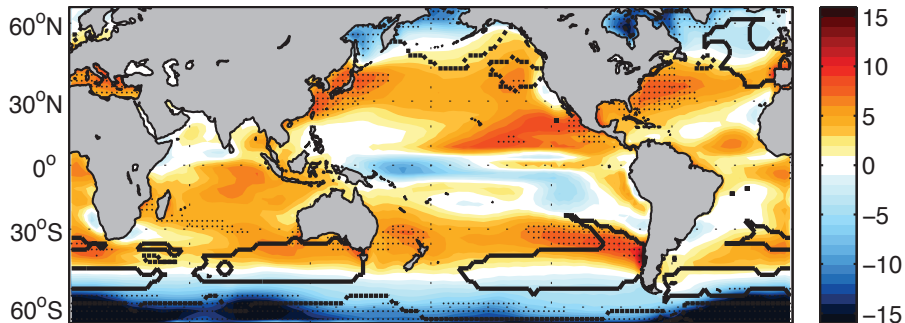
Qu et al. (2015) reach similar conclusion for subtropical clouds.

Implications for Climate Feedbacks

Response to 4xCO₂ Forcing (CMIP5 Multi-Model Mean)

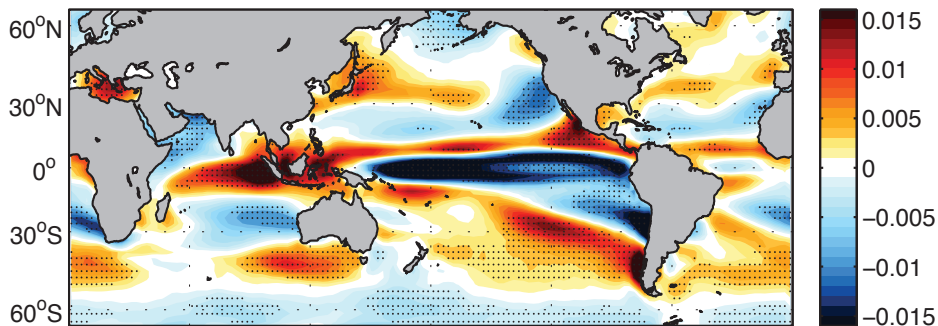
Shortwave CRE

W/m²



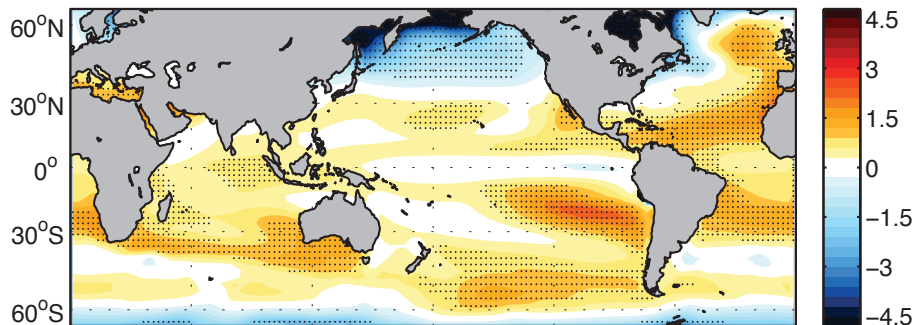
ω_{500}

Pa s⁻¹



EIS

K



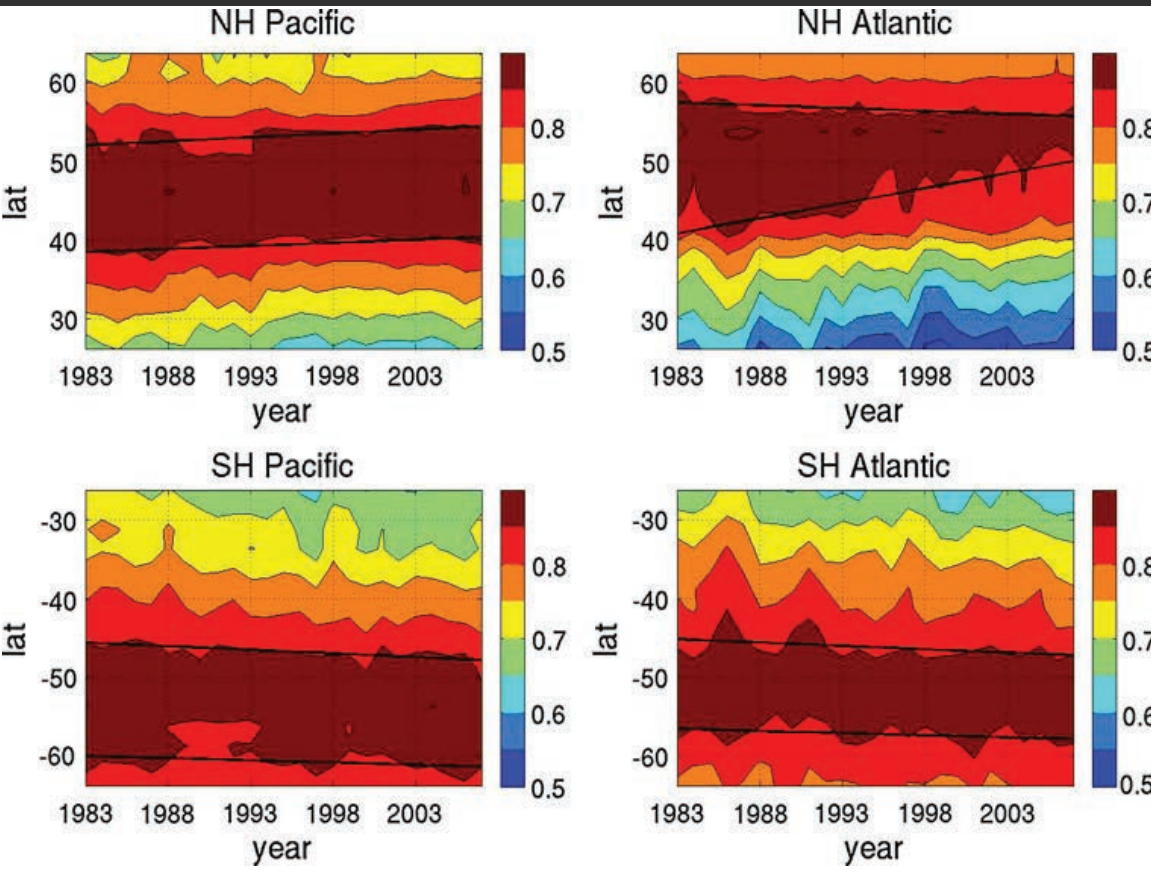
In response to increasing greenhouse gases, many climate models project similarly signed changes in EIS and vertical velocity over mid-latitude oceans.

Should we have lower confidence in dynamically driven cloud feedbacks in type I models in these regions?

Evidence from Observed Trends

Bender et al. (2012)

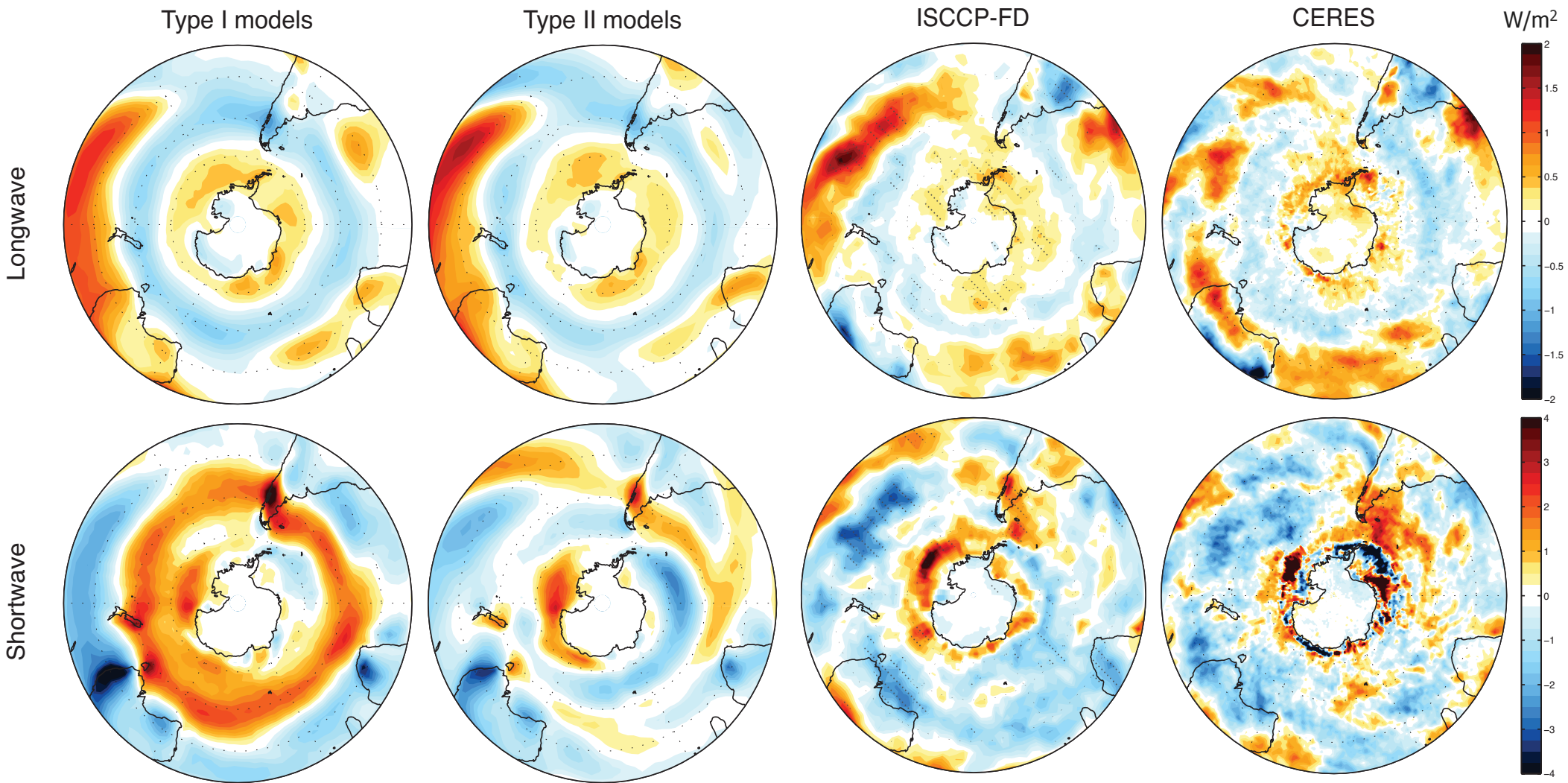
Annual-mean Total Cloud Fraction



- ISCCP satellite observations suggest poleward shift in mid-latitude cloud patterns.
 - Reduction in total cloud cover
 - Increase in high cloud cover
 - Positive net radiative effect
- Poleward shift in mid-latitude cloud bands also apparent in surface-based cloud observations (Eastman and Warren 2013).

CMIP5 Models

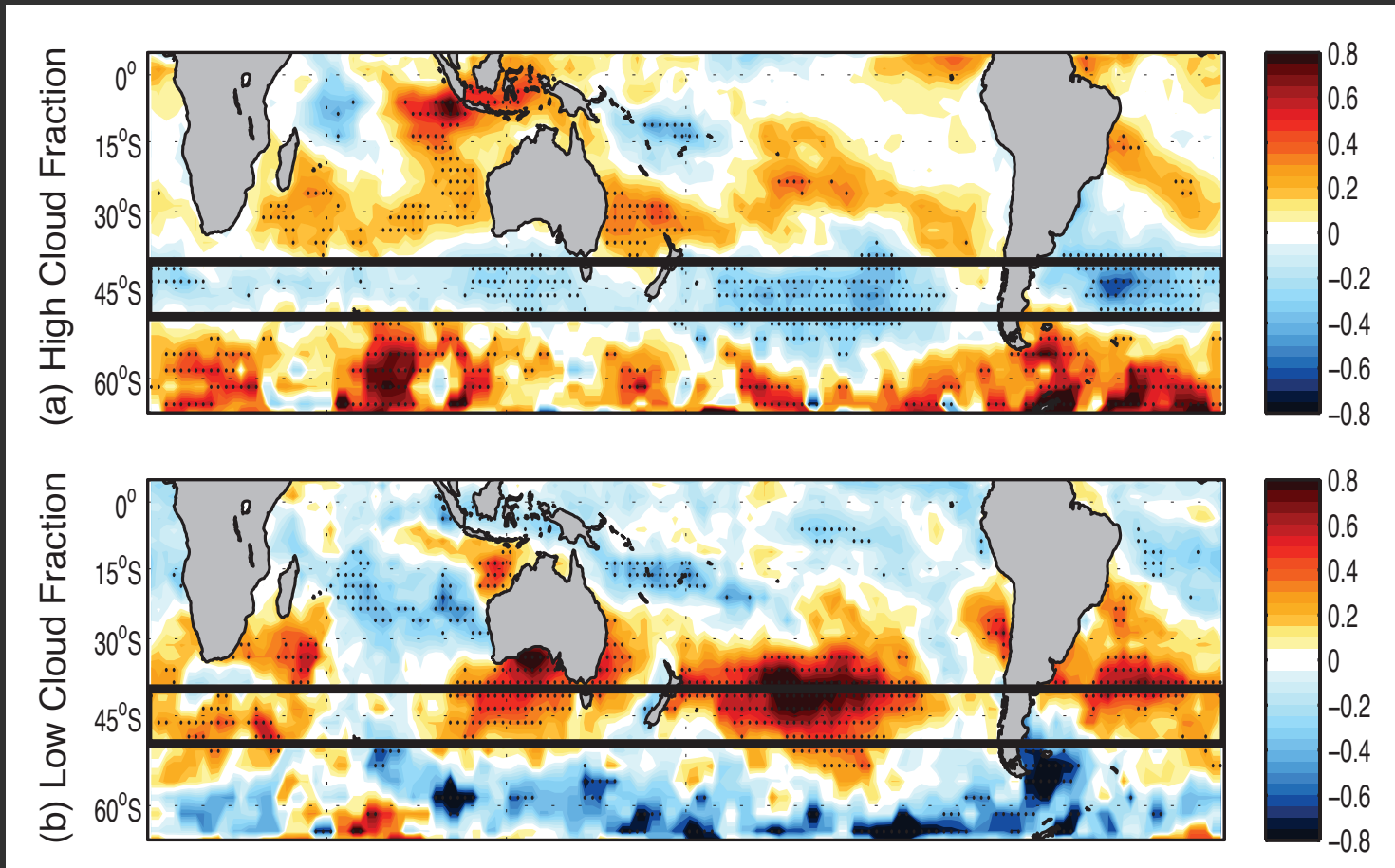
Satellite Observations



Cloud-Radiative Effect Anomalies
Associated with 1° Poleward Jet Shift

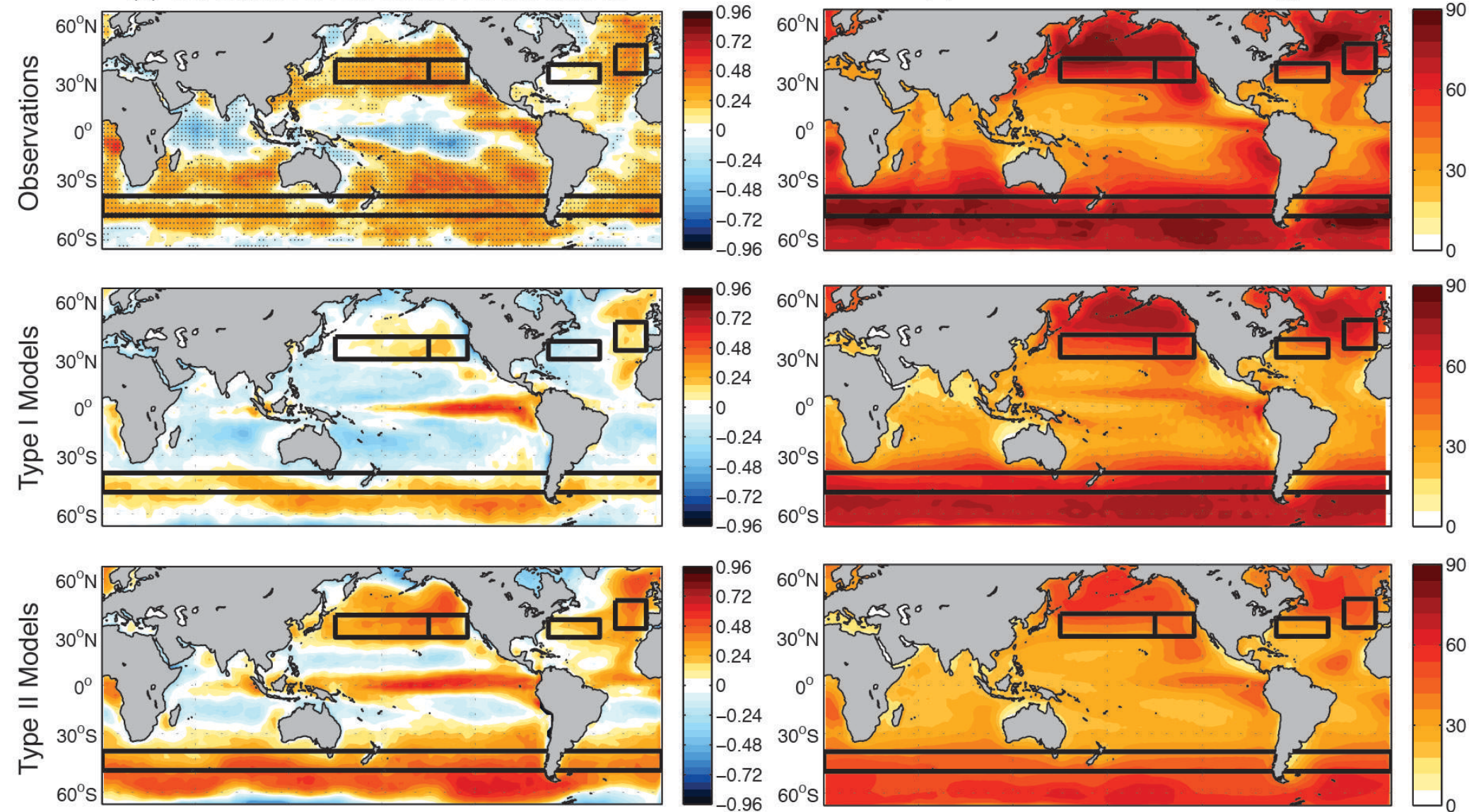
Cloud Anomalies

1° Poleward Jet Shift (Observed Variability)



(a) Correlations of Low Cloud Fraction with EIS

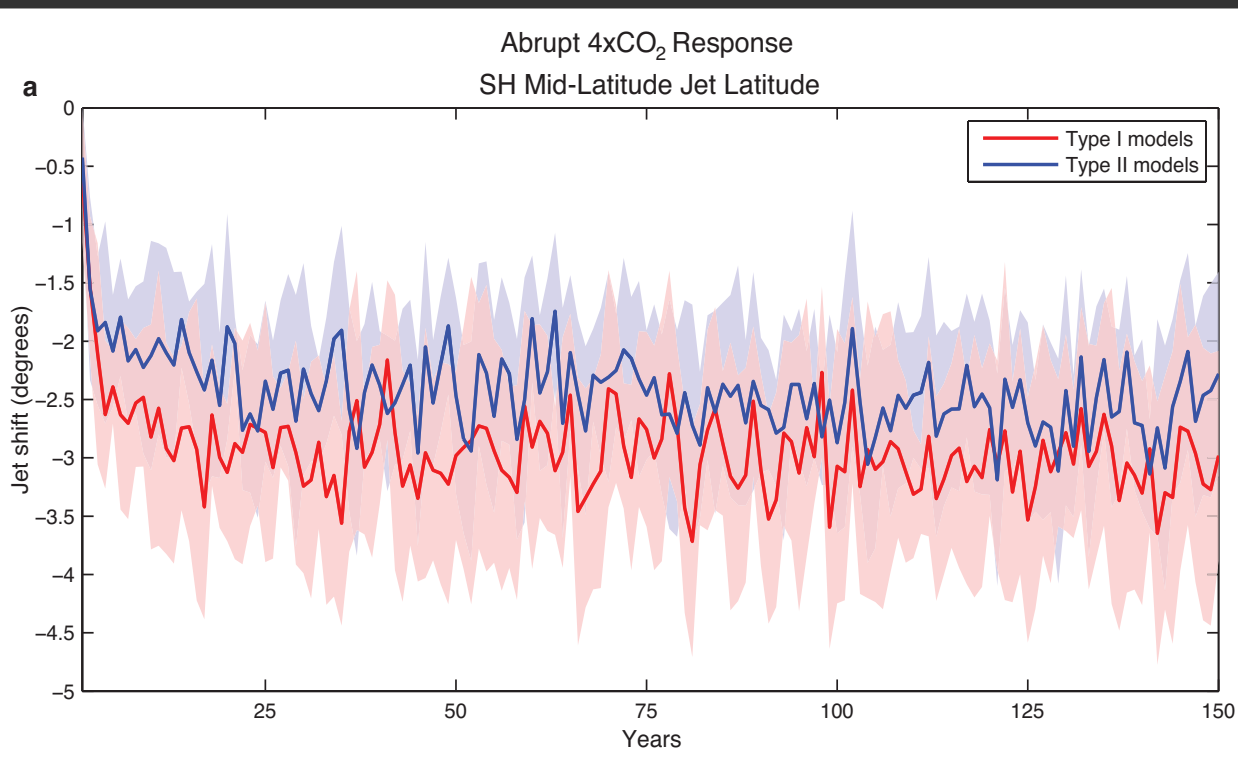
(b) Low Cloud Fraction Climatology



Implications for Climate Change

Idealized Experiment: Abrupt Quadrupling of CO₂

- 1) Jet shifts rapidly poleward in both classes of CMIP5 models.

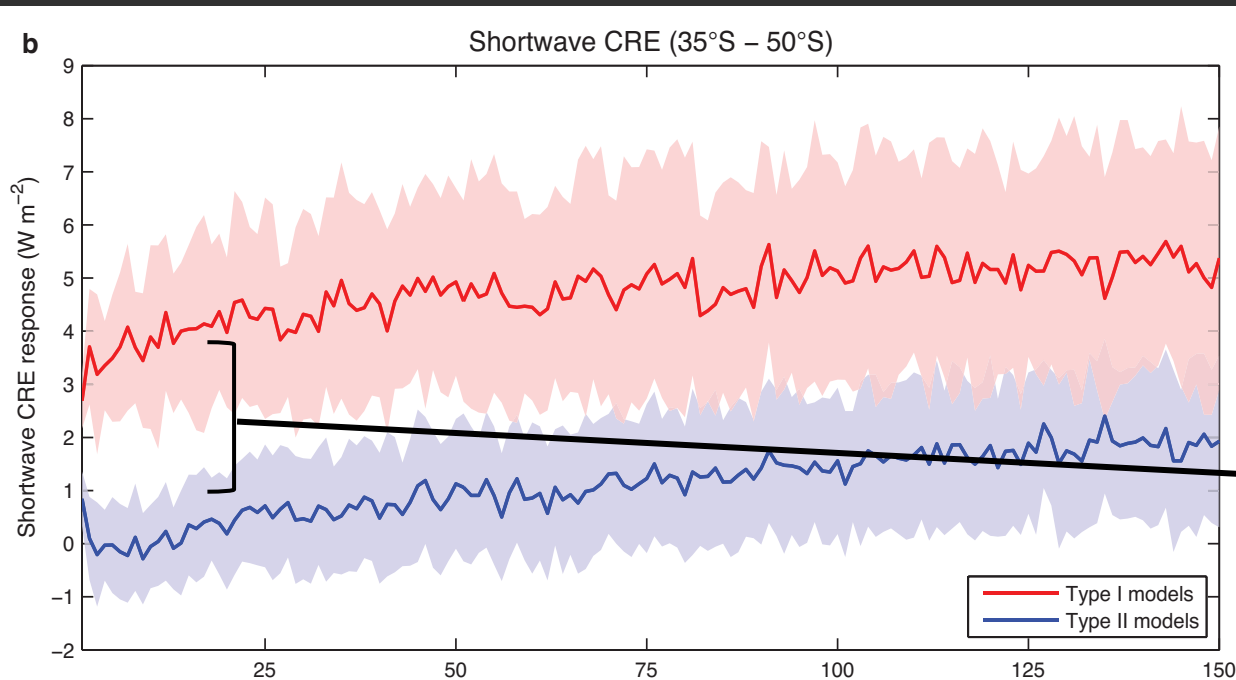


Grise and Polvani (2014, J. Climate)

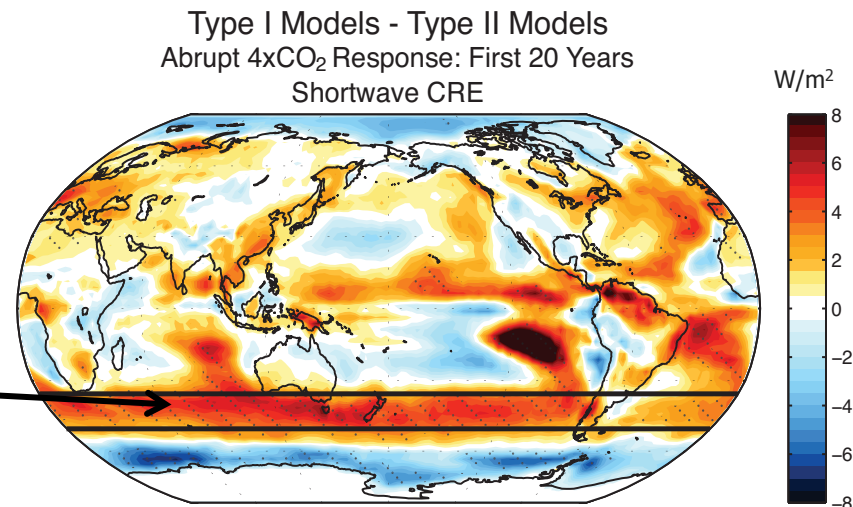
Implications for Climate Change

Idealized Experiment: Abrupt Quadrupling of CO₂

- 1) Jet shifts rapidly poleward in both classes of CMIP5 models.
- 2) Rapid reduction in reflection of shortwave radiation by clouds in Type I models only



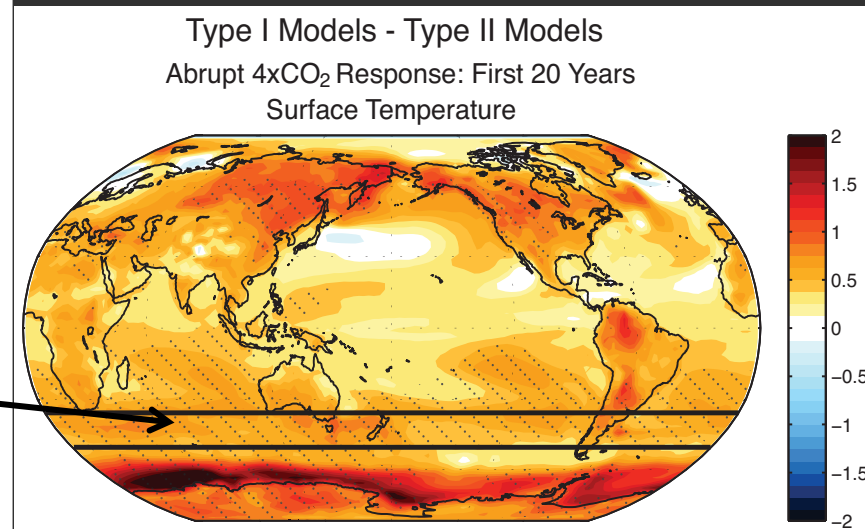
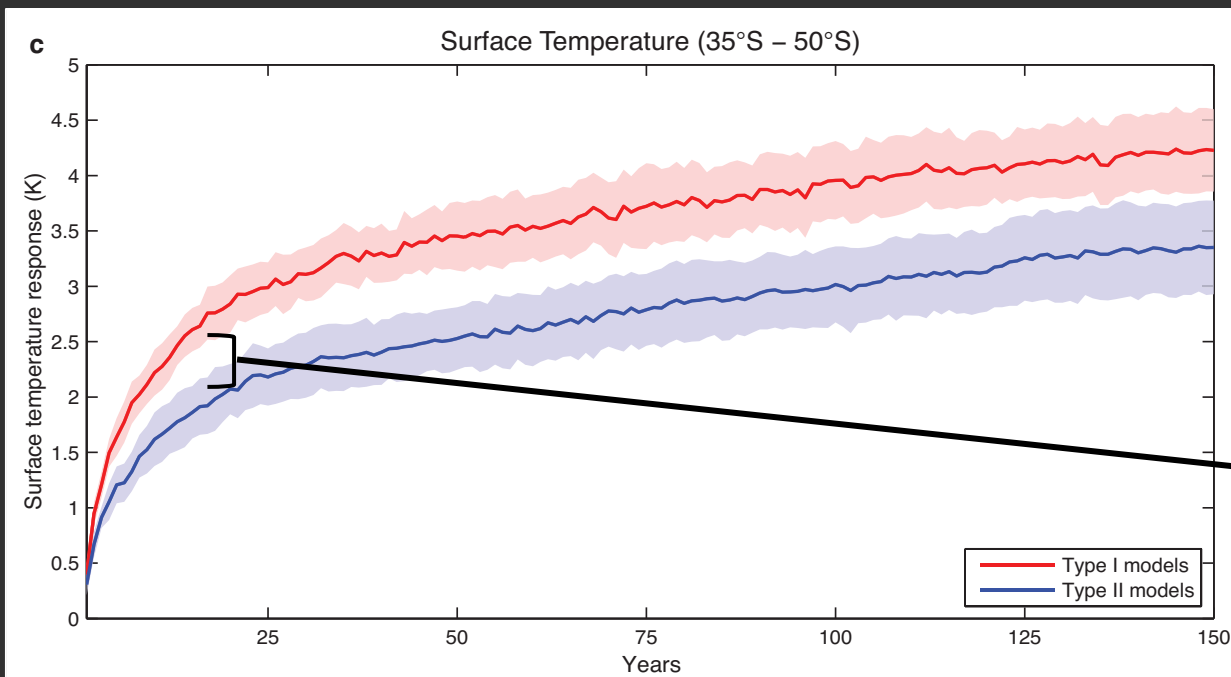
Grise and Polvani (2014, J. Climate)



Implications for Climate Change

Idealized Experiment: Abrupt Quadrupling of CO₂

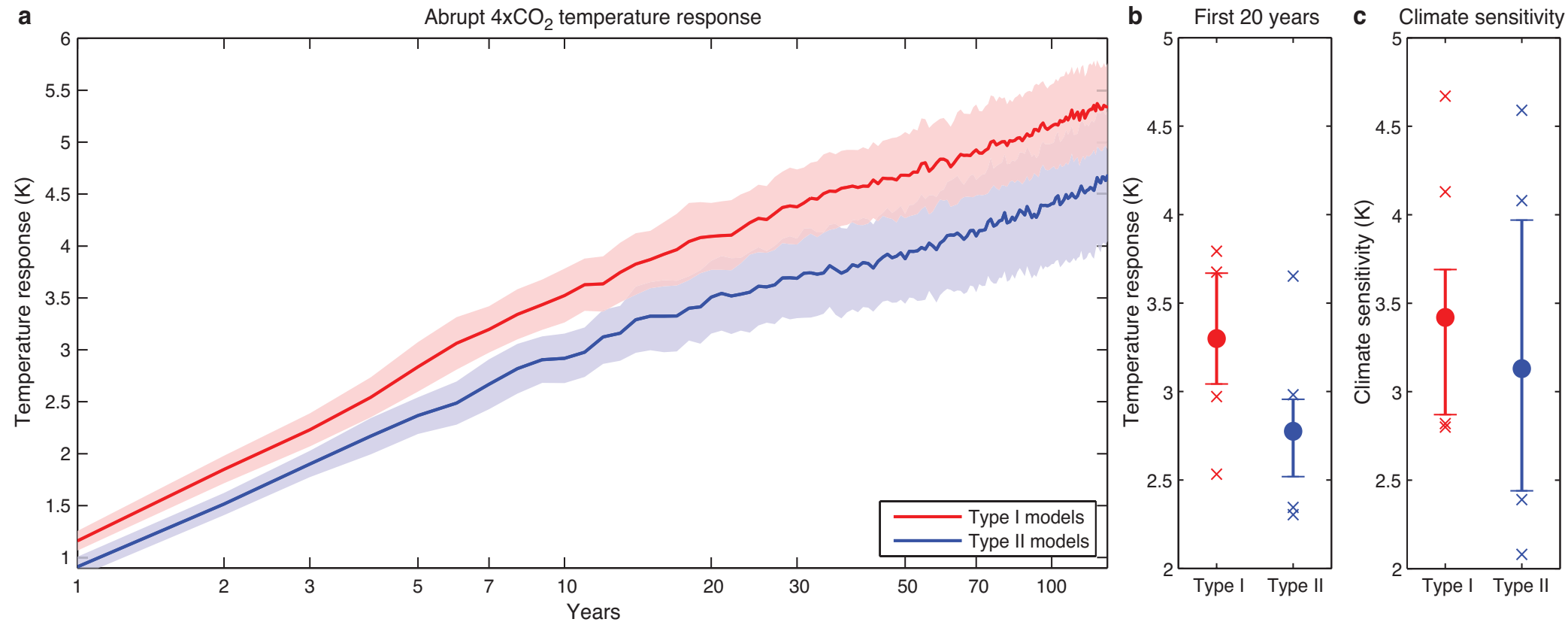
- 1) Jet shifts rapidly poleward in both classes of CMIP5 models.
- 2) Rapid reduction in reflection of shortwave radiation by clouds in Type I models only
- 3) Excess initial warming in SH in Type I models



Grise and Polvani (2014, J. Climate)

Implications for Climate Change

Idealized Experiment: Abrupt Quadrupling of CO₂

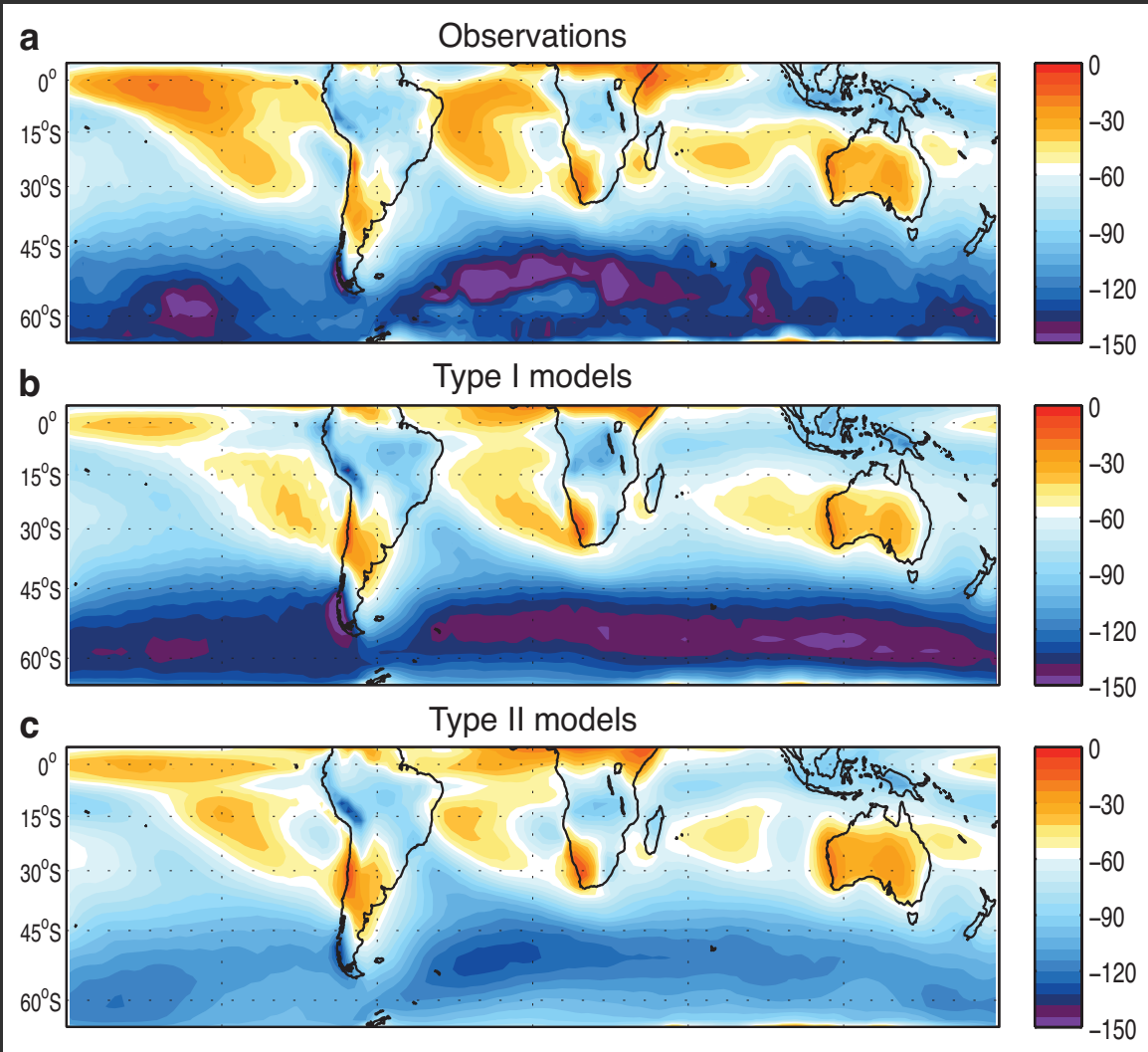


Grise and Polvani (2014, J. Climate)

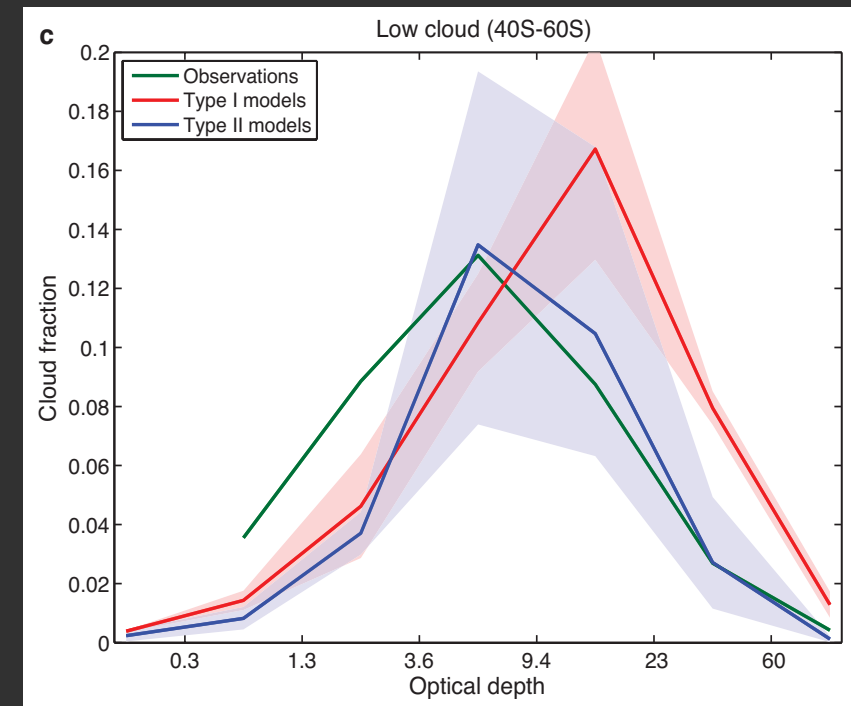
Nearly instantaneous global warming response is significantly larger in the type I models.

Why do the models behave so differently?

Shortwave Cloud-Radiative Effect Climatology



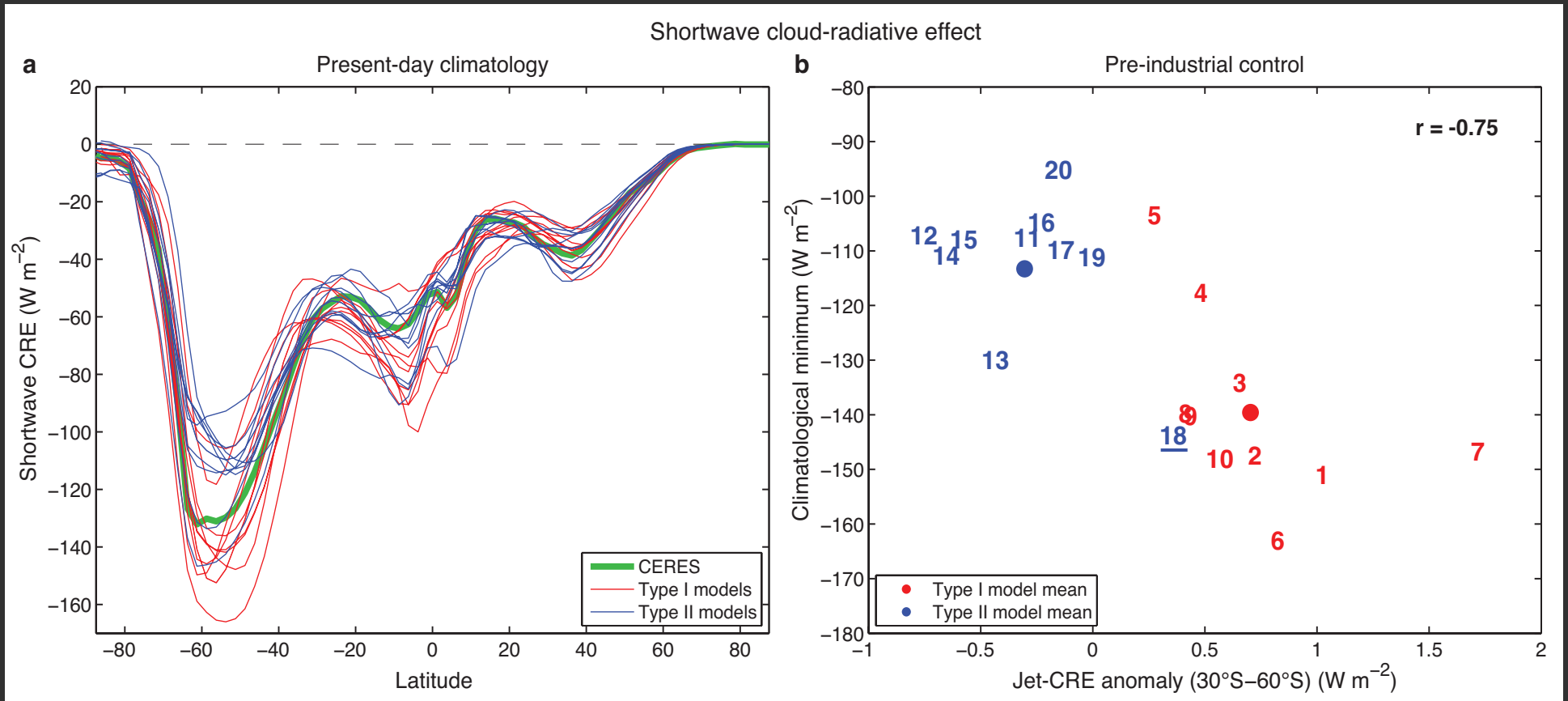
Cloud Fraction Climatology



Grise and Polvani (submitted)

- Type I models: Bright, zonally symmetric Southern Ocean clouds
- Type II models: Less bright, more zonally asymmetric clouds

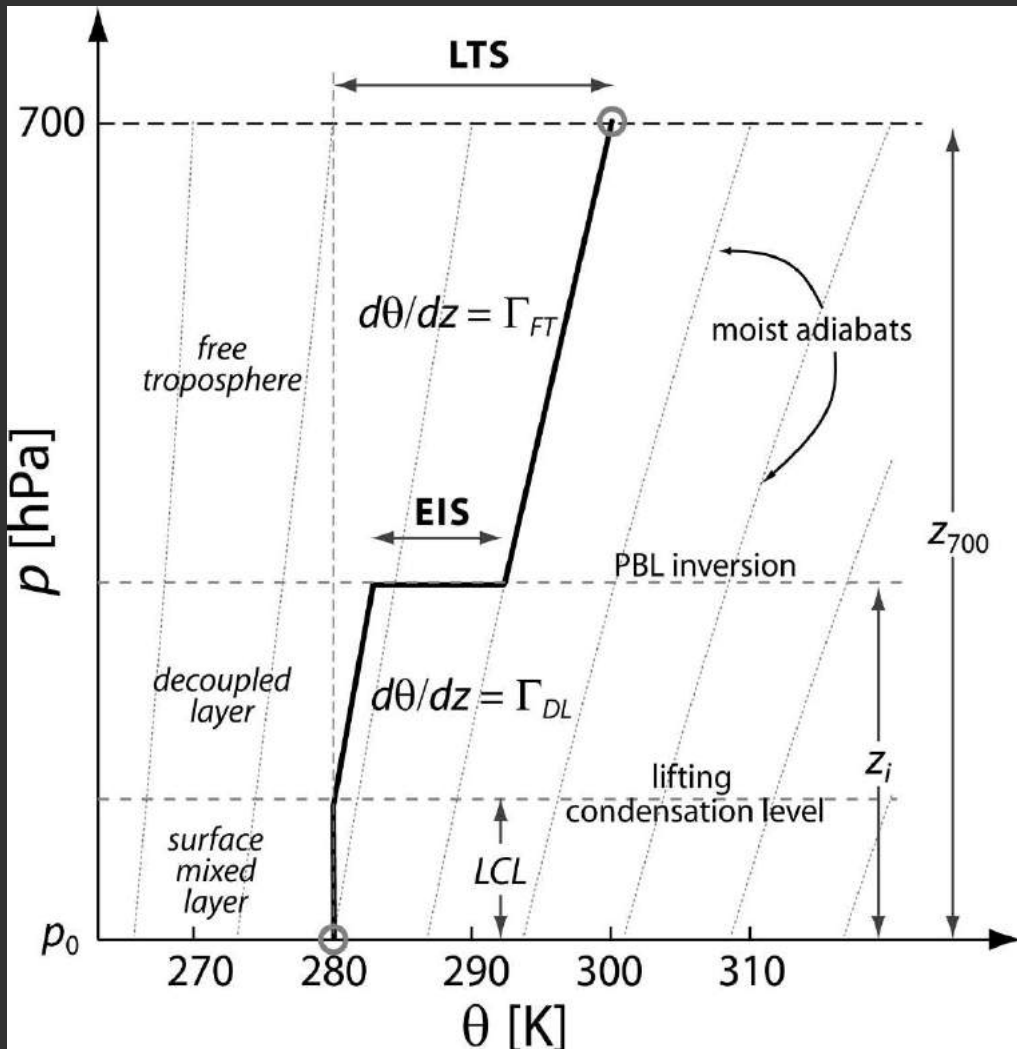
CMIP5 CRE Zonal Mean Climatology



Estimated Inversion Strength (EIS)

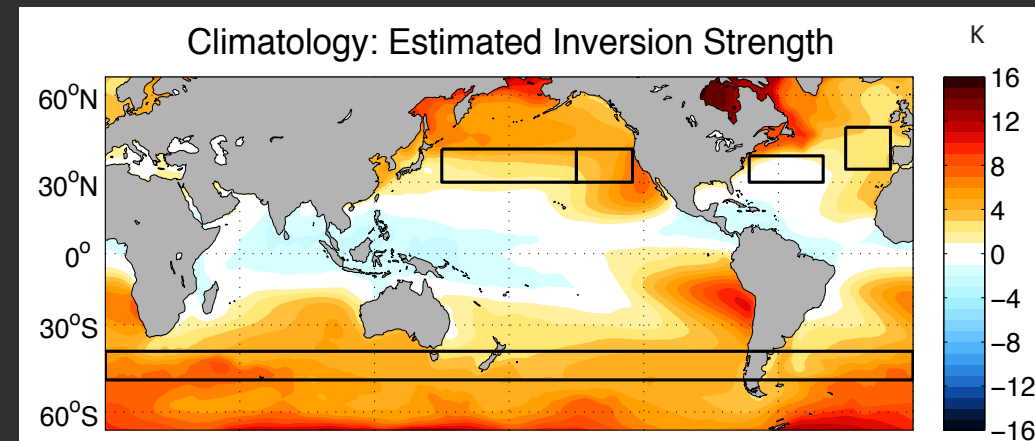
Wood and Bretherton (2006)

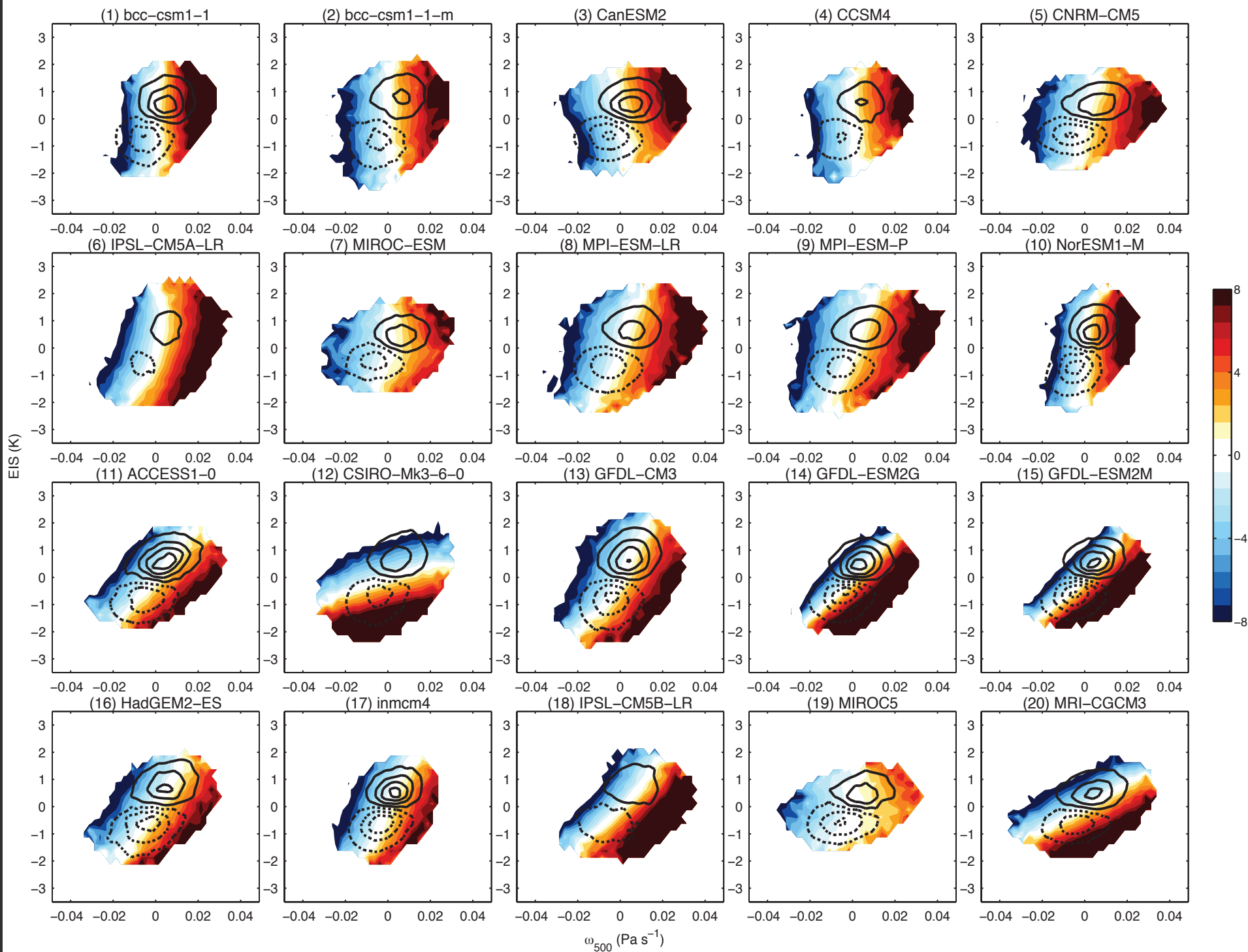
- $$EIS = \theta_{700 \text{ hPa}} - \theta_{\text{sfc}} - \Gamma_{\text{moist}}(850 \text{ hPa}) (z_{700 \text{ hPa}} - z_{\text{LCL}})$$



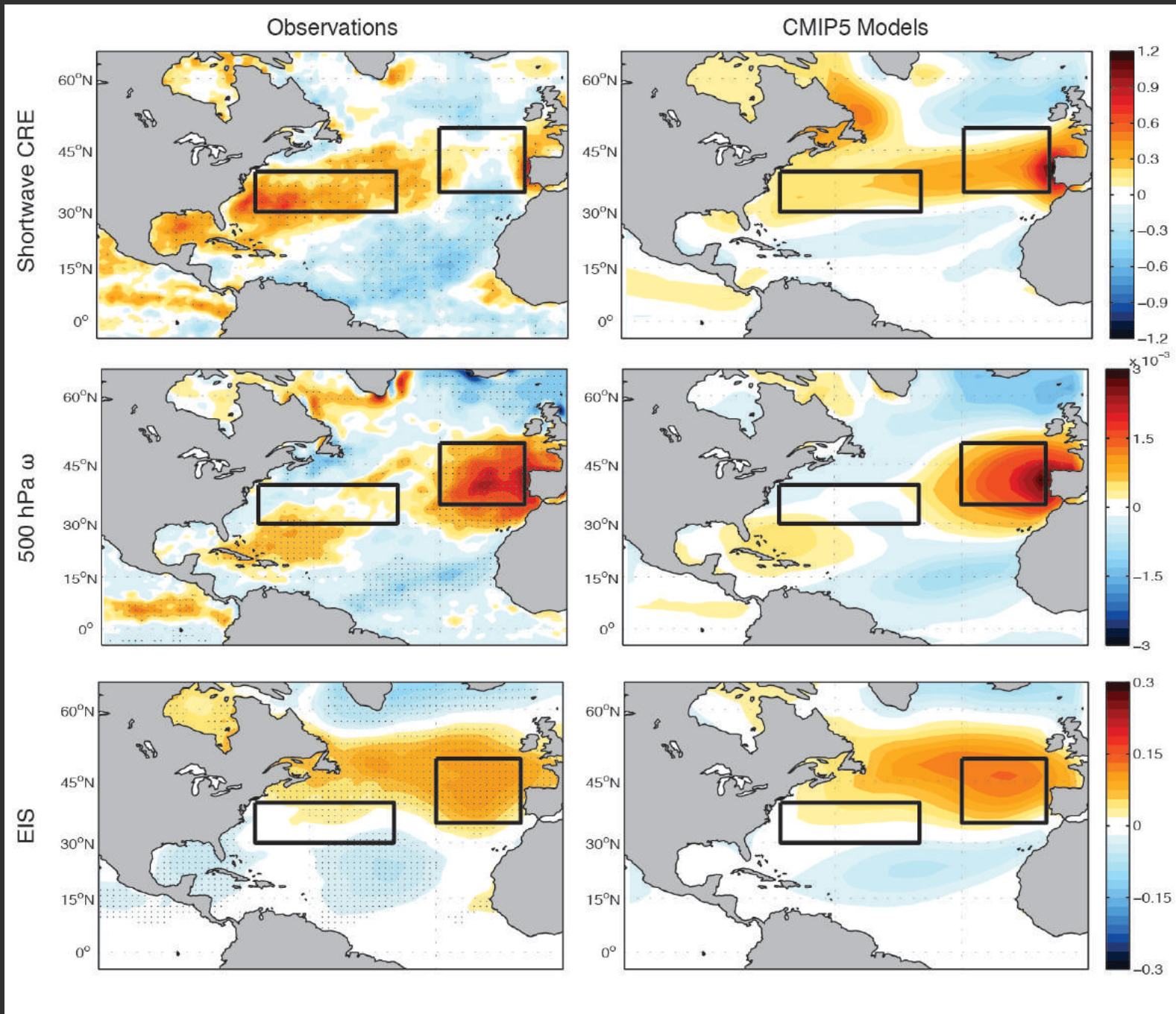
$$\Delta\theta = \theta_{700 \text{ hPa}} - \theta_{\text{sfc}} - \Gamma_{\text{FT}} (z_{700} - z_i) - \Gamma_{\text{DL}} (z_i - \text{LCL})$$

$$\Delta\theta = \theta_{700 \text{ hPa}} - \theta_{\text{sfc}} - \Gamma_{\text{FT}} z_{700} + \Gamma_{\text{DL}} \text{LCL}$$

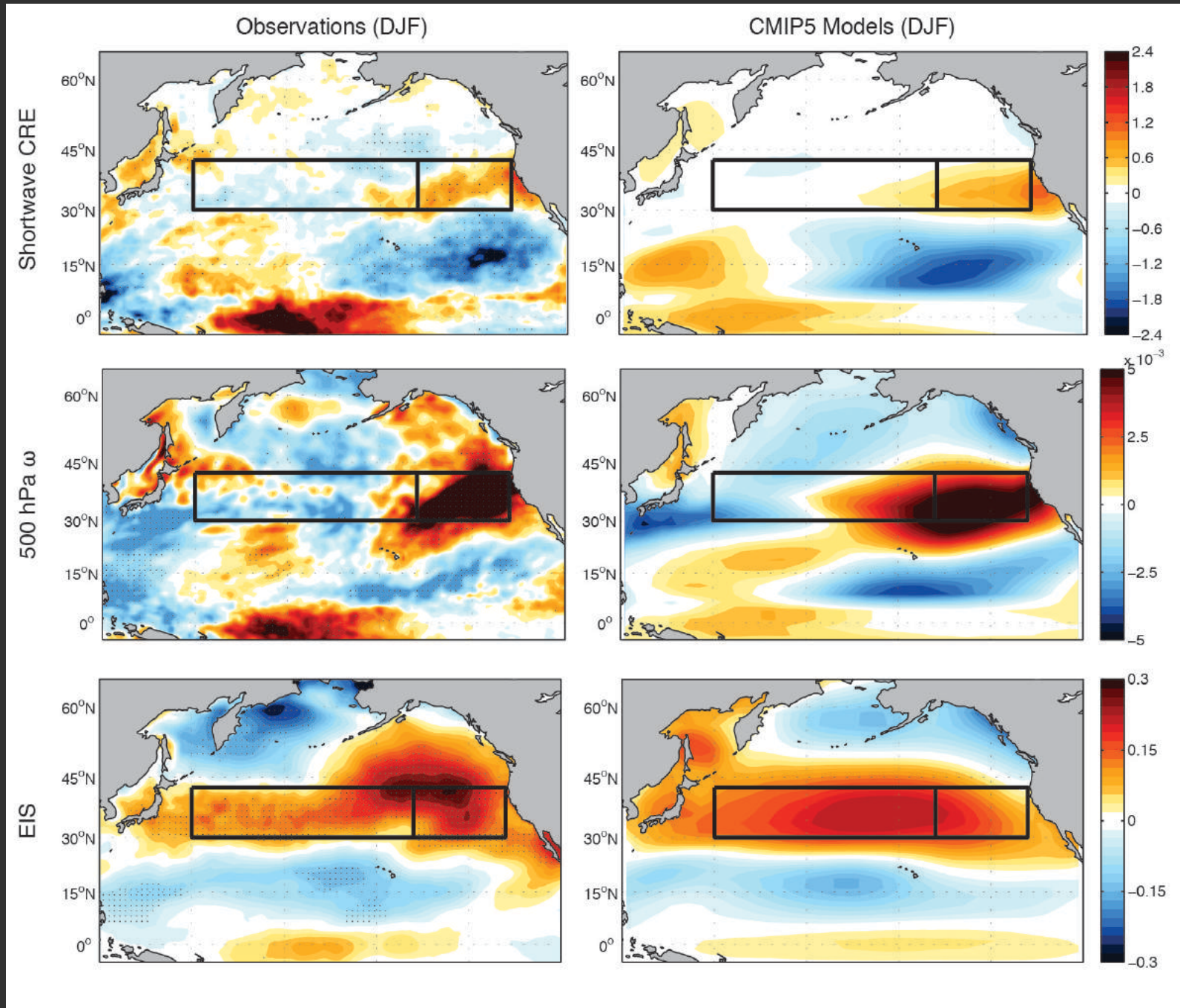




What about the Northern Hemisphere?



What about the Northern Hemisphere?



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